

Seasonal change in soil nitrogen mineralization in young *Chamaecyparis obtusa* stands at the upper and lower positions on a slope in central Japan

Nanae Hosokawa ⁽¹⁻²⁾,
Sasa Tajima ⁽²⁾,
Hajime Kobayashi ⁽³⁾,
Keizo Hirai ⁽⁴⁾

Nitrogen (N) is a critical element for tree growth in forest ecosystems. As trees absorb inorganic N, the soil N mineralization process is a key process for their development. Although the spatial pattern of soil N mineralization is expected to relate to tree growth, the difference between the upper and lower positions within a small-scale slope is unclear. Therefore, we compared annual and seasonal soil N mineralization rates in Japanese cypress (*Chamaecyparis obtusa* [Siebold & Zucc.] Endl.), which stands at both the upper and lower positions on a slope. We used the resin-core method to estimate *in situ* soil N mineralization rates. Additionally, the litter decomposition rate and inorganic N passed through the litter layer, which are primary sources for soil N mineralization, were investigated using the litter bag and resin-core methods. Our findings revealed that the annual soil N mineralization rate at the lower position was 5 times higher than that at the upper position. Moreover, seasonal variations in soil N mineralization rate tended to be higher at the lower position than at the upper position. The temporal change in input ammonium passed through the litter layer was similar to that of the nitrification rate in the soil at the lower position, except for winter. Notably, high nitrification in winter at the lower position may be related to soil frost, which can accelerate the decomposition of organic matter. Despite these differences, the litter decomposition rate was similar between the slope positions. The higher soil N mineralization rate and substrate input may result in higher tree growth at the lower position on a slope.

Keywords: Soil Nitrogen Mineralization, Field Soil Incubation, *Chamaecyparis Obtusa*, Resin Core Method

Introduction

Nitrogen (N) and soil water are limiting factors for tree growth (Girardin et al. 2016, Du et al. 2020, Nakayama et al. 2024). The availability of N in the soil is considered a significant indicator of tree growth (Kawada 1989). Although the spatial pattern of soil N mineralization rate is expected to be linked to tree growth, their relationship within a small-scale slope is unclear.

In situ soil incubation methods aim to cap-

ture the effect of natural environmental fluctuations on mineralization (Eno 1960, DiStefano & Gholtz 1986, Hanselman et al. 2004). Generally, the rate of soil N mineralization varies with soil temperature, moisture, and substrate quantity and quality (Tokuchi et al. 2014). Since the moisture environment drastically changes on a small scale on a slope (Shigyo et al. 2022), capturing moisture fluctuations is necessary. The resin-core method places soil between

ion-exchange resins, which allow water input into the soil and output from the soil without N contamination (DiStefano & Gholtz 1986). It can also evaluate the amount of substrates input into the soil by the upper resin.

Here, we evaluate annual and seasonal changes in soil N mineralization rate using the resin-core method at the upper and lower positions on a slope. Additionally, the litter decomposition rate and nitrogen input into the soil were evaluated.

Methods

Study site

The study was conducted at the Terasawa-yama station (Education and Research Center of Alpine Field Science, Shinshu University, Nagano prefecture, Japan). The mean annual air temperature and precipitation during 2006-2010 were 9.0 °C and 1479.4 mm, respectively (Kinoshita & Suzuki 2010, Kinoshita et al. 2011).

We established plots on a slope in a young Japanese cypress (*Chamaecyparis obtusa*, hereafter hinoki) forest (planted in 1985 – Tab. 1). The study site is located at an altitude of 1080 to 1140 m a.s.l., in the east-northeastern direction, with an aver-

□ (1) Shikoku Research Center, Forestry and Forest Products Research Institute, 2-915 Asakuranishi, Kochi-shi, Kochi, 780-8077 (Japan); (2) Division of Forest and Environmental Symbiosis Science, Faculty of Agriculture, Shinshu University, Nagano 399-4511 (Japan); (3) Education and Research Center of Alpine Field Science, Shinshu University, Nagano 399-4511 (Japan); (4) Department of Forest Soils, Forestry and Forest Products Research Institute, Matsunosato 1, Tsukuba, Ibaraki 305-8687 (Japan)

@ Nanae Hosokawa (hosokawa_nanae660@ffpri.go.jp)

Received: Aug 19, 2024 - Accepted: May 06, 2025

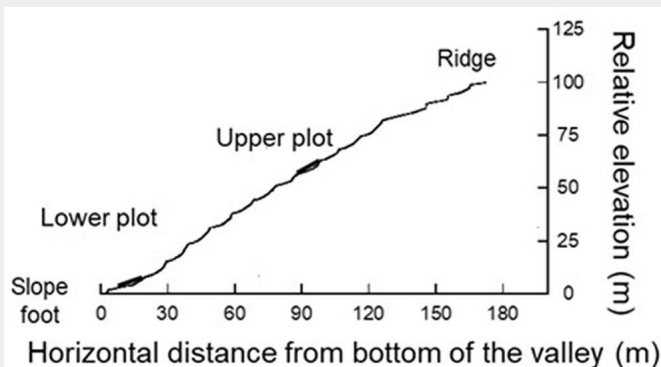
Citation: Hosokawa N, Tajima S, Kobayashi H, Hirai K (2025). Seasonal change in soil nitrogen mineralization in young *Chamaecyparis obtusa* stands at the upper and lower positions on a slope in central Japan. iForest 18: 197-201. - doi: 10.3832/ifor4710-018 [online 2025-07-20]

Communicated by: Claudia Cocozza

Tab. 1 – Hinoki cypress growth at the upper and lower plots. Values indicate mean \pm standard deviation. The p-value indicates the result of the t-test. The number of trees measured was 42 and 47 in the upper and the lower plot, respectively.

Parameter	Upper plot	Lower plot	p-value
Height (m)	9.3 \pm 0.9	11.7 \pm 2.8	< 0.05
Diameter of breast height (cm)	11.7 \pm 1.5	12.8 \pm 2.0	< 0.05
D ² H (m ³)	0.13 \pm 0.04	0.20 \pm 0.08	< 0.05

Fig. 1 – Location of the plots along a small-scale slope in the Hinoki cypress study stand. The figure is based on Hosokawa et al. (2012).



age slope angle of 35°. The bedrock is weathered granite. The hinoki stands were planted at a density of 3,500 trees ha⁻¹ and maintained at 2,600-2,900 trees ha⁻¹ after non-commercial cutting.

The plot (height 8 m and width 20 m) was established at the middle and bottom of the slope (hereafter upper and lower plots, respectively). The horizontal separation between the two plots is 80 m (Fig. 1). According to the Japanese local soil classification, the soil is classified as semidried Brown forest soil (BD_(d)) and moderate Brown forest soil (BD – Kubo 1982), respectively. The BD soil corresponds to Cambisols in the international classification (IUSS Working Group WRB 2022). Litter layer was found on the upper plot but not on the lower plot. The soil in the upper plot was poorer and more acidic, with lower soil N content and pH as well as a higher C/N ratio, compared to the soil in the lower plot (Tab. 2).

Soil environment

The soil temperature was measured at 5 and 25 cm depths using a thermometer (TR-5220®, T & D Ltd., Matsumoto, Japan). The soil water potential at depths of 20

and 35 cm was measured using a tension meter (KDC-S5, Kona System Ltd., Sapporo, Japan) in 2010.

Soil N mineralization rate

The resin core method (DiStefano & Gholtz 1986) was used to measure soil N mineralization and the inorganic N passing through the litter layer (hereafter referred to as input N). First, a 100 cm³ polyvinyl chloride column (diameter 2.5 cm, height 5 cm) was used to collect mineral soil from three depths (0-5, 20-25, and 50-55 cm). Second, an ion exchange column was attached to the upper and lower sides of the column. Finally, the columns were buried back in the ground. The ion exchange resin was a mixture of 20 g of cation exchange resin (Amberlite IR-120B, Organo, Tokyo, Japan) and anion exchange resin (Amberlite IRA-410, Organo, Tokyo, Japan). The 0-5 cm resin core was buried to the same depth as the soil surface and covered with the litter layer. The resin cores measuring 20-25 cm and 50-55 cm were buried at the same depth from which they were collected. The soil incubation was conducted from 2011 to 2012. The 0-5 and 20-25 cm resin cores were used to detect seasonal

change and annual soil N mineralization. The 50-55 cm resin core was used to estimate annual soil N mineralization. Each incubation had three replications.

When the incubation finished, the soils were passed through a 2 mm sieve to remove coarse detritus and gravel. The soil and the ion exchange resin were extracted by 2 M KCl and 1M KCl, respectively. The extract was frozen until the chemical analysis was completed. A flow injection analyzer was used to test ammonium and nitrate (Quatro 2HR®, BL-Tech, Osaka, Japan).

The mineralization in soil N during a given period (mg N kg⁻¹ period⁻¹) was calculated as follows (eqn. 1):

$$N_m = S_a + R_a - S_i \quad (1)$$

where N_m is the N mineralization over time, S_a is the inorganic N in the soil column at the end of the incubation, R_a is the inorganic N in the bottom resin column at the end of the incubation, and S_i is the inorganic N in the soil at the start of the incubation. The soil N mineralization rate (mg N kg⁻¹ day⁻¹) was the ratio of soil N during a given period and its incubation days.

Litter decomposition

In November 2010, we collected fresh leaf litter at the upper and lower plots. For the litter decomposition experiment, 6 g of air-dried hinoki leaf litter was placed in a square bag (20 × 20 cm). The litter collected from the upper plot was placed on the upper plot, and the litter collected from the lower plot was placed on the lower plot. The initial carbon (C) and N contents in the leaf litter from the upper plot were significantly lower than that from the lower plot, and the C/N ratio of leaf litter from the upper plot was significantly higher than that from the lower plot (t-test – total C: 62.2 \pm 0.14% and 62.6 \pm 0.15%; total N: 0.52 \pm 0.01% and 0.76 \pm 0.01%; C/N ratio: 120 \pm 1.79 and 82.4 \pm 0.83, respectively). In February 2011, the litter bag was placed on the plot and collected after 79, 161, 224, 317, 444 days. The litter samples were dried for 48 hours at 75 °C, and weighed. The dry combustion method was used to determine total C and N. The ash contents of the litter were determined by measuring ignition loss at 880 °C, and the organic matter and amount of C or N were calculated as follows (eqn. 2 to eqn. 4):

$$M_o = M_T - M_R \quad (2)$$

$$M_C = M_o \cdot C_C \quad (3)$$

$$M_N = M_o \cdot C_N \quad (4)$$

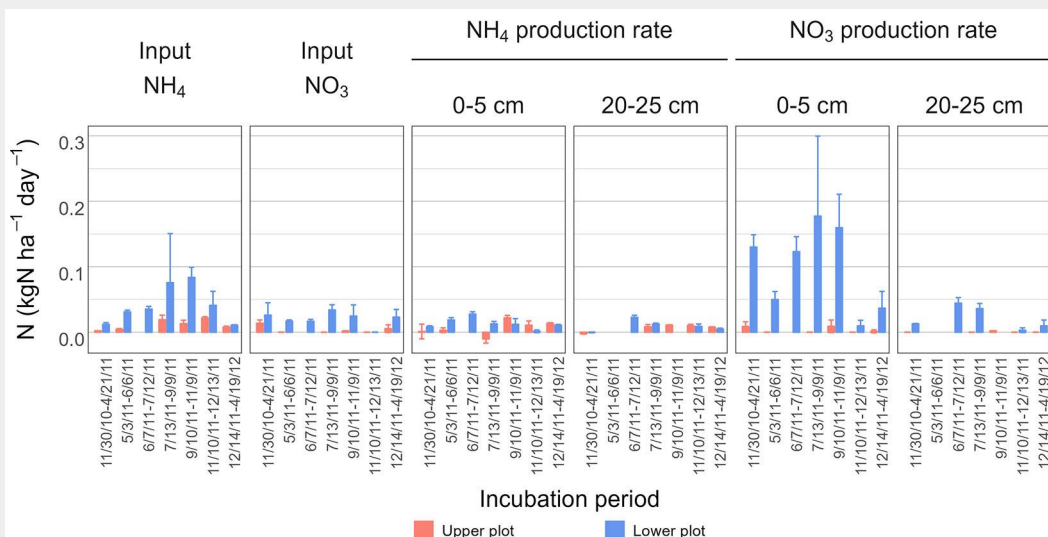
where M_o is the organic matter content, M_T is the total litter content, M_R is the residue content (including ash and dust), M_C and M_N are carbon and nitrogen contents, respectively, C_C and C_N are carbon and nitrogen concentration, respectively.

Because a significant amount of soil ad-

Tab. 2 – Soil chemical properties at the two plots. Values indicate the mean \pm standard deviation. Data from Hosokawa et al. (2012).

Plot position	Depth (cm)	pH (H ₂ O)	pH (KCl)	C (gC kg ⁻¹)	N (gN kg ⁻¹)	C/N ratio
Upper plot	0-5	4.68 \pm 0.02	3.61 \pm 0.02	48.3	3.3	14.6
	20-25	5.05 \pm 0.02	4.18 \pm 0.03	40.0	2.6	15.4
Lower plot	0-5	5.30 \pm 0.09	4.05 \pm 0.02	43.5	4.1	10.6
	20-25	5.56 \pm 0.08	4.42 \pm 0.02	32.1	3.1	10.4

Fig. 2 - Seasonal changes in input N and soil N mineralization rate. Input NH_4 and Input NO_3 denote ammonium and nitrate passed through the litter layer, respectively. Vertical bars represent the standard deviation ($n = 3$). The date is displayed in Month/Day/Year format.



hered to the litter samples from the lower plot, we calculated the organic matter content by subtracting the ash from the total weight.

Data analysis

Welch's *t*-test was used to compare the annual/seasonal soil N mineralization rates and the litter decomposition rates between plots. We considered multiplicity in tests of seasonal soil N mineralization and litter decomposition using the Benjamini & Hochberg (BH) method (Benjamini & Hochberg 1995). The significant *p*-value was set as < 0.05 . All statistical analyses and graphics were conducted using R version v. 4.1.2 (R Core Team 2021).

Results and discussion

Soil temperature at the upper plot was higher than in the lower plot (Fig. S1 in Supplementary material), and the annual mean of the soil temperature at 5 and 25 cm depths in the upper plot was 2 °C higher

than in the lower plot. The temperature difference could be due to the light conditions on the slope. Indeed, the lower plot, located near the valley, is shaded by the opposing slope and is covered by a canopy of Japanese cypress trees. In contrast, the upper plot receives shade later and is primarily dominated by deciduous broadleaf trees.

The soil water potential at 20 and 35 cm depths at the upper plots was around -10 kPa (Fig. S2 in Supplementary material), except for the period July-September. The annual mean soil water potential at 20 and 35 cm depths in the upper plot was -13 and -11 kPa, respectively, whereas it was -9 and -8 kPa in the lower plot. This indicates that the lower plot retained more moisture than the upper plot. The maximum soil water potential at 20 cm depth in the upper plot was -50 kPa, while it was only -25 kPa in the lower plot, indicating that during the summer the upper plot experienced more severe soil desiccation than the lower plot.

Input NH_4 tended to be higher than input NO_3 (Fig. 2). Both input NH_4 and NO_3 at the lower plot tended to be higher than those at the upper plot, and significantly higher at some points (Tab. 3). Annual input NH_4 at the lower plot was significantly higher than that at the upper plot (*t*-test, $p < 0.05$), although there is no difference in NO_3 between the two slope positions (Fig. 3). The amount of Japanese cypress leaf litter that remained one year later was comparable to previous studies (Takeda 1995, Makita & Fujii 2015). The N levels remaining in litter samples from both plots decreased to 79 days before increasing again (Fig. 4d), showing no significant difference between the plots at other times (Tab. 4). The litter immobilized N rather than facilitating the mineralization during the experiment (Fig. 4). As a result, litter decomposition was not associated with the input N. One year after the experiment began, there were no significant differences in the remaining organic matter and N between

Tab. 3 - Results of the multiple *t*-tests of input N and soil N mineralization rates following the BH method. The date is shown in the Year/Month/Day format. Asterisks indicate a significant difference between the upper and lower plots. (NA): inorganic N not detected; (+): data not available; (++): data in the upper plot not available; (+++): data in the lower plot not available.

Periods	Input NH_4	Input NO_3	Soil NH_4 production rate		Soil NO_3 production rate	
			0-5 cm	20-25 cm	0-5 cm	20-25 cm
2010/11/30-2011/04/21	*	Not significant	Not significant	*	*	*
2011/05/03-2011/06/06	*	*	*	+	*	+
2011/06/07-2011/07/12	++	++	++	++	++	++
2011/07/13-2011/09/09	Not significant	*	*	Not significant	Not significant	*
2011/09/10-2011/11/09	*	Not significant	Not significant	+++	Not significant	+++
2011/11/10-2011/12/13	Not significant	NA	Not significant	Not significant	Not significant	NA
2011/12/14-2012/04/19	Not significant	Not significant	Not significant	Not significant	Not significant	Not significant

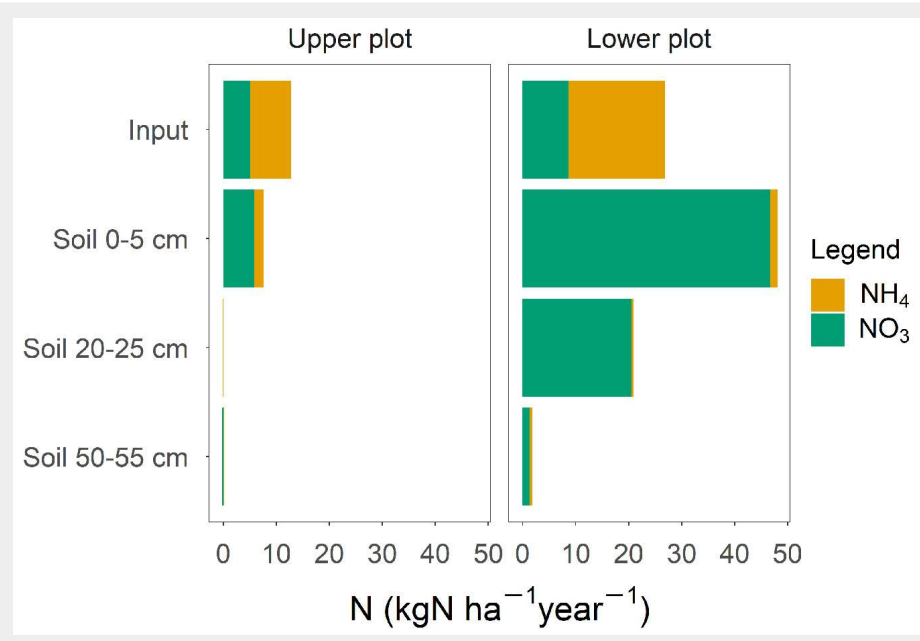


Fig. 3 - The average annual rate of input N and soil N mineralization. Input, Soil 0-5 cm, Soil 20-25 cm, and Soil 50-55 cm indicate inorganic N passed through the litter layer, and soil N mineralization at 0-5, 20-25, and 50-55 cm (n = 3-5).

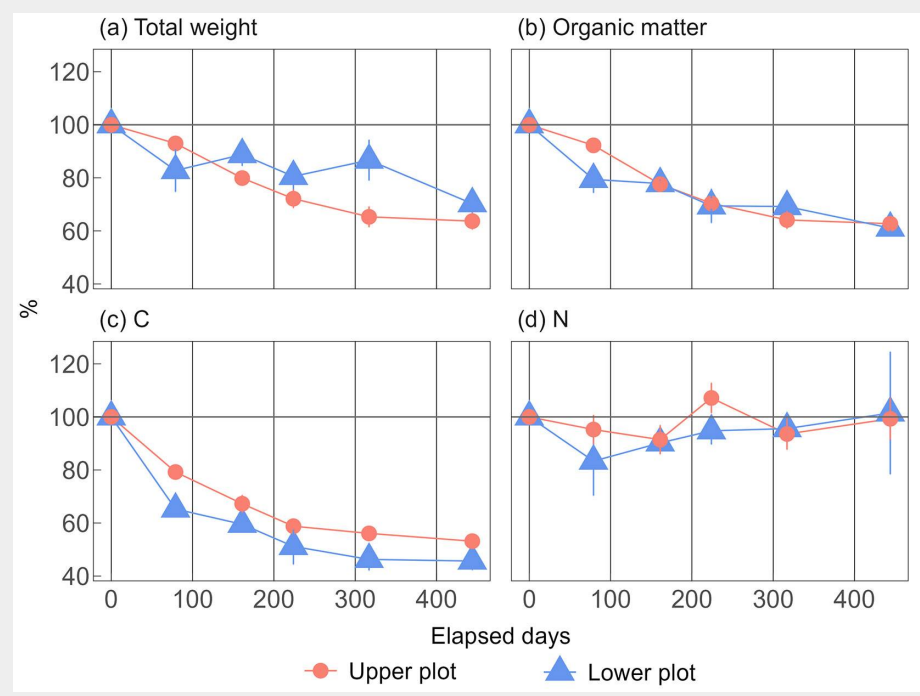


Fig. 4 - Temporal changes in the Hinoki cypress leaf litter. (a): Total weight; (b): organic matter; (c): total C; (d): total N. Vertical bar at each point represents the standard deviation (n = 3-5).

the upper and the lower plots (Tab. 4). The variation in annual nitrogen mineralization in the two plots closely reflected the seasonal changes (Fig. 3). When combining input N with annual soil NH_4 and NO_3 production, the sum at the lower plot was significantly higher than at the upper plot (t-test, $p < 0.05$). Indeed, the sum of NH_4 and NO_3 in the lower plot was five times greater than in the upper plot.

Tab. 4 - Results of the multiple t-tests of remaining total weight, organic matter, C and N, following the BH method. Asterisks indicate a significant difference between the upper and lower plots. (NA): not available.

Elapsed days	Total weight	Organic matter	C	N
0	NA	NA	NA	NA
79	Not significant	Not significant	*	Not significant
161	*	Not significant	*	Not significant
224	Not significant	Not significant	Not significant	Not significant
317	*	Not significant	*	Not significant
444	*	Not significant	*	Not significant

An unexpected result was found in the annual changes in soil N mineralization rate during winter. Given the low temperatures, we anticipated a low soil N mineralization rate during winter. Contrary to our expectations, we observed a surprisingly high rate of soil N mineralization during this season. This finding is consistent with a recent report that the winter soil N mineralization accounted for 10-30% of the annual amount in Japan (Hirai et al. 2007). High nitrification in winter at the lower position may be related to soil frost, which can accelerate the decomposition of organic matter (Hosokawa et al. 2017). Further, we found that the soil N mineralization rate in the 2011-2012 winter was notably lower than that in the 2010-2011 winter. Hishi et al. (2014) found that soil freeze-thaw had a negative impact on nitrification rates in northern

Japan. The mean soil temperature during incubation in 2010-2011 and 2011-2012 was -0.55°C and -0.72°C , respectively, indicating that the 2011-2012 winter was colder than the previous winter. This implies that the soil frost in 2011-2012 was greater than that in 2010-2011, resulting in lower soil N mineralization in the 2011-2012 winter. This variation in soil N mineralization during winter may contribute to differences in the N cycle between the plot positions.

Conclusion

We investigated soil N mineralization at the upper and lower positions along a small-scale slope in a hinoki cypress (*Chamaecyparis obtusa*) stand in Japan. The soil N mineralization at the lower plot was significantly higher than that at the upper plot. Although N input into the soil at the lower plot was significantly higher than that at the upper plot, the litter decomposition rate was similar between the plots, as the litter immobilized N during the experiment. An unexpected finding was observed in the annual changes in the soil N mineralization rate during winter. The annual changes in soil frost, soil water content, and litter fall may have contributed to this annual change in soil N mineralization.

Acknowledgments

NH carried out the field survey and laboratory measurements, and wrote the manuscript. ST carried out laboratory measurements. KH and HK conceived the study and helped draft the manuscript. We appreciate Drs. Hosoo, Ono, and Oda for their assistance with soil analysis. We also appreciate the members of the Forest Environment and Ecology Laboratory at Shinshu University for their assistance. Two anonymous reviewers provided valuable comments. This study was partly supported by KAKENHI 23K13988.

Declaration of interest statement

We have no competing interests.

References

Benjamini Y, Hochberg Y (1995). Controlling the false discovery rate: a practical and powerful approach to multiple testing. *Journal of the Royal Statistical Society, series B* 57: 289-300. - doi: [10.1111/j.2517-6161.1995.tb02031.x](https://doi.org/10.1111/j.2517-6161.1995.tb02031.x)
DiStefano JF, Gholtz HL (1986). A proposed use of ion exchange resin to measure nitrogen min-

eralization and nitrification in intact soil cores. *Communications in Soil Science and Plant Analysis* 17: 989-998. - doi: [10.1080/00103628609367767](https://doi.org/10.1080/00103628609367767)

Du E, Terrer C, Pellegrini AFA, Ahlström A, Van Lissa CJ, Zhao X, Xia N, Wu X, Jackson RB (2020). Global patterns of terrestrial nitrogen and phosphorus limitation. *Nature Geoscience* 13: 221-226. - doi: [10.1038/s41561-019-0530-4](https://doi.org/10.1038/s41561-019-0530-4)

Eno CF (1960). Nitrate production in the field by incubating the soil in polyethylene bags. *Soil Science Society of America Proceedings* 24: 277-279. - doi: [10.2136/sssaj1960.03615995002400040019x](https://doi.org/10.2136/sssaj1960.03615995002400040019x)

Girardin MP, Hogg EH, Bernier PY, Kurz WA, Guo XJ, Cyr G (2016). Negative impacts of high temperatures on growth of black spruce forests intensify with the anticipated climate warming. *Global Change Biology* 22: 627-643. - doi: [10.1111/gcb.13072](https://doi.org/10.1111/gcb.13072)

Hanselman TA, Graetz DA, Obreza TA (2004). A comparison of *in situ* methods for measuring net nitrogen mineralization rates of organic soil amendments. *Journal of Environmental Quality* 33: 1098-1105. - doi: [10.2134/jeq2004.1098](https://doi.org/10.2134/jeq2004.1098)

Hirai K, Noguchi K, Mizoguchi T, Kaneko S, Takahashi M (2007). Contribution of subsoil and season for nitrogen mineralization under field conditions in forest soil. *Japanese Journal of Forest Environment* 49: 51-59. [in Japanese with English summary]

Hishi T, Urakawa R, Tashiro N, Maeda Y, Shibata H (2014). Seasonality of factors controlling N mineralization rates among slope positions and aspects in cool-temperate deciduous natural forests and larch plantations. *Biology and Fertility of Soils* 50: 343-356. - doi: [10.1007/s00374-013-0863-x](https://doi.org/10.1007/s00374-013-0863-x)

Hosokawa N, Ito D, Kobayashi H, Hirai K (2012). Annual change of soil nitrogen mineralization in a young Hinoki cypress stand (*Chamaecyparis obtusa*) at middle and lower slope positions. *Bulletin of Shinshu University AFC* 10: 109-114. [in Japanese with English summary]

Hosokawa N, Isobe K, Urakawa R, Tateno R, Watanabe T, Fukuzawa K, Shibata H (2017). Soil freeze-thaw with root litter alters N transformations during the dormant season in soils under two temperate forests in northern Japan. *Soil Biology and Biochemistry* 114: 270-278. - doi: [10.1016/j.soilbio.2017.07.025](https://doi.org/10.1016/j.soilbio.2017.07.025)

IUSS Working Group WRB (2022). World reference base for soil resources. International soil classification system for naming soils and creating legends for soil maps (4th edn). International Union of Soil Sciences - IUSS, Vienna, Austria.

Kawada H (1989). Introduction to forest soil science. Hakuyu-sha, Tokyo, Japan, pp. 223-239. [in Japanese]

Kinoshita Y, Suzuki J (2010). Meteorological data of Terasawa-yama Station, AFC, Shinshu University, 2009. *Bulletin of Shinshu University AFC* 8: 85-86. [in Japanese]

Kinoshita Y, Suzuki J, Kobayashi H (2011). Meteorological data of Terasawa-yama Station, AFC, Shinshu University, 2010. *Bulletin of Shinshu University AFC* 9: 123-124. [in Japanese]

Kubo T (1982). Classification of forest soils in Japan. *JARQ: Japan Agricultural Research Quarterly* 16: 212-217. [online] URL: https://www.jircas.go.jp/sites/default/files/publication/jarq/16-3-212-217_o.pdf

Makita N, Fujii S (2015). Tree species effects on microbial respiration from decomposing leaf and fine root litter. *Soil Biology and Biochemistry* 88: 39-47. - doi: [10.1016/j.soilbio.2015.05.005](https://doi.org/10.1016/j.soilbio.2015.05.005)

Nakayama M, Abe Y, Atarashi-Andoh M, Tange T, Sawada H, Liang N, Koarashi J (2024). Quantitative importance of subsoil nitrogen cycling processes in Andosols and Cambisols under temperate forests. *Applied Soil Ecology* 201: 105485. - doi: [10.1016/j.apsoil.2024.105485](https://doi.org/10.1016/j.apsoil.2024.105485)

R Core Team (2021). R: A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. [online] URL: <http://www.R-project.org/>

Shigyo N, Furusawa H, Yamashita N, Nagakura J, Manaka T, Yamada T, Hirai K (2022). Slope-induced factors shape bacterial communities in surface soils in a forested headwater catchment. *Catena* 214: 106253. - doi: [10.1016/j.catena.2022.106253](https://doi.org/10.1016/j.catena.2022.106253)

Takeda H (1995). A 5 year study of litter decomposition processes in a *Chamaecyparis obtusa* Endl. forest. *Ecological Research* 10: 95-104. - doi: [10.1007/BF02347659](https://doi.org/10.1007/BF02347659)

Tokuchi N, Yoneda S, Ohte N, Usui N, Koba K, Kuroiwa M, Toda H, Suwa Y (2014). Seasonal changes and controlling factors of gross N transformation in an evergreen plantation forest in central Japan. *Journal of Forest Research* 19: 77-85. - doi: [10.1007/s10310-013-0391-4](https://doi.org/10.1007/s10310-013-0391-4)

Supplementary Material

Fig. S1 - Soil temperature at the upper and the lower plots.

Fig. S2 - Soil water potential at the upper and the lower plots.

Link: Hosokawa_4710@suppl001.pdf