

Dynamics of soil organic carbon (SOC) content in stands of Norway spruce (*Picea abies*) in central Europe

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Norway spruce is the main forest tree species in the Czech Republic. Until now, little attention has been given in the literature to the dynamics of soil organic carbon (SOC) content under Norway spruce stands as a function of stand characteristics. The aim of this study is to estimate soil organic carbon (SOC) content and stock changes in organic and surface mineral soil horizons on forest sites with a dominant representation of Norway spruce. In the study area, a significantly higher content of SOC was found in the surface mineral soil horizon than in the organic soil horizon. In both soil horizons, there was evidence of an increasing trend of SOC with the increasing age of forest stands, a decreasing trend of SOC with increasing density of stocking and an increasing trend of SOC with increasing altitude. The relationship of SOC content with soil group (Podzol vs. non-Podzol) has also been demonstrated. The greatest potential for long-term carbon sequestration in soils was shown in older stands (101-190 years) dominated by Norway spruce with lower density of stocking, located in forest vegetation zones (altitude range: 1010-1225 m a.s.l.) where natural mountain Norway spruce forests currently occur. According to our results, Norway spruce stands may perform a stable function of carbon sequestration in the soil at these sites, especially in the mineral soil horizon.

Keywords: Forest Soil, C Sequestration, *Picea abies*, Site Conditions, Stand Characteristics, Climate Change

Introduction

It is estimated that forest ecosystems currently store more than 80% of all terrestrial aboveground carbon (C) and approximately 45-70% of all soil carbon (Post et al. 1990, Birdsey et al. 1993, Lorenz & Lal 2010). The role of forests in the global carbon cycle, and in mitigation of the impact of climate change, is clear and fairly well documented (IPCC 2000, Janssens et al. 2003, Ciais et al. 2008, Lorenz & Lal 2010). The study of the carbon cycle and its stocks and fluxes has come to the forefront of interest over the last 10 years, especially in relation to limiting the effects of global warming.

In terms of the predicted impact of climate change (the scenario of "Business as Usual" developed by IPCC and the global

scenarios by Goddard Institute – Čermák et al. 2004) and according to regional scenarios for the Czech Republic (in 2030), it is likely that the most affected stands will be the Norway spruce (*Picea abies* [L.] Karst.) stands in the current 3rd and 4th forest vegetation zones (FVZ – Viewegh et al. 2003), where it is expected that they will almost completely disappear. The occurrence of Norway spruce in this area will likely be restricted to the highest mountainous border region (from the 5th FVZ upwards – Čermák et al. 2004, Čermák & Holuša 2010, Hlásny et al. 2011, Fitzgerald & Lindner 2013, Hanewinkel et al. 2013).

Norway spruce represents about 38% of trees in the European forests, and it is the dominant forest tree species (52-53%) in the Czech Republic, although its share

should not exceed 11% in the natural forest tree communities of this region. Indeed, Norway spruce naturally occurs on modal sites from the 5th FVZ (520-730 m a.s.l.), with an admixture (up 20%) of beech (*Fagus sylvatica*) and silver fir (*Abies alba*), up to 8th FVZ (1170-1225 m a.s.l.) where it is the dominant tree species. However, Norway spruce in the Czech Republic is also cultivated outside of its optimum elevation range (in particular, in the 2nd, 3rd and 4th FVZ), and this could be one of the factors related to the significant dieback of this tree species in Czech forests in recent decades.

According to several authors (Berger et al. 2002, Körner 2006) Norway spruce growing in its ecological optimum will play a relatively important role in facing global climate change, mainly in terms of carbon sequestration and storage. However, it is very important to know the optimum conditions for carbon sequestration, as the potential for sequestration of carbon is influenced by site conditions (such as climate, temperature, etc. – Wiesmeier et al. 2013, Tashi et al. 2016) and soil conditions (soil moisture, pH, nutrients, texture), as well as species composition, the applied form of management and its intensity (Jandl et al. 2007, Vesterdal et al. 2008, Wiesmeier et al. 2013, Vilén et al. 2015, Lundmark et al. 2016).

The aims of this study were: (i) to analyse the state and development of the carbon stock in the organic (O) and surface min-

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eral (A) soil horizon in forest sites where Norway spruce is the dominant species; (ii) to study the relationships of selected stand characteristics (age, density of stocking and absolute height-yield class) with SOC content in the organic and surface mineral soil horizons under different site conditions (altitude, soil type, humus form); (iii) to assess the ability of Norway spruce stands to sequester carbon during stand development and their response to productive forest management in relation to vegetation zonality; (iv) to estimate optimum conditions for carbon sequestration in Norway spruce stands. The study could contribute to better understand current dieback of Norway spruce forests in central Europe and under the expected climate change.

Material and methods

Study area

For the purpose of this study, data was taken from the database of forest typology compiled by the Czech Forest Management Institute (FMI, Brandýs nad Labem, Czech Republic – <http://www.uhul.cz/what-we-do/forest-typology>). The database includes information about forests from monitoring plots established all over the country, including soil conditions, species composition and mensurational characteristics of forest stands.

Forty-two monitoring plots were selected from the forest typology database, including their soil samples taken in the years 1975 to 2003. The size of the selected circular plots varied in the range of 400-500 m². All the selected plots are located in the eastern part of the Czech Republic, in the Jeseníky Mountains and in the Drahanská vrchovina Upland (Fig. S1 in Supplementary material), at elevation ranging from 300 to 1318 m a.s.l., with mean annual air temperature (MAAT) varying from 2.1 to 8.1 °C and mean annual precipitation (MAP) between 617 mm and 1106 mm.

Soil conditions at the selected plots are variable. According to the World Reference Base (WRB 2014), soil groups were mainly represented by Cambisols, Podzols (Haplic and Entic Podzol), Stagnosols and Retisols. The dominant soil group in the selected set are Cambisols, and the prevailing humus form is moder. In each soil pit, two soil horizons (organic and surface mineral) were differentiated. The thickness of organic (O) horizons in the studied plots ranged from 2 to 15 cm (mean = 6.55 cm), while the thickness of surface mineral (A) horizons ranged from 1 to 27 cm (mean = 8.80 cm). The depth of the two horizons (O+A horizons) was in the range of 5-30 cm (mean = 15.33 cm).

Stand description

We selected monitoring plots with dominant representation (91-100%) of Norway spruce. Forest sites at the monitoring plots were included in units of the Typological

forest classification system (Viewegh et al. 2003) used in the forests of the Czech Republic. This system includes site conditions (climate, soil, nutrients) and their relation to forest stands, including phytocenosis. In the Typological forest classification system (and in forest typology) the term of “forest vegetation zone” (FVZ) is used. FVZ is a formalized forestry unit expressing the relationship between climate and biocenosis, and describes in a simplified form the vegetation zones. According to the above classification system, the selected plots were included to FVZ from the 2nd to the 9th. Samples from the plots were further classified in ecological series and edaphic categories (see Tab. S1 in Supplementary material).

The monitoring plots mostly consist of productive forest under continuous management. Species composition is regulated in the first 10 years by the reduction of density of advance growth and thinning out seedlings. From 10 to 25 years of age, tree species composition and its spatial distribution is regulated through cleaning (juvenile thinning). At age 25-60, thinning is carried out to increase the ecological stability and resilience, and to support the growth of promising individuals. At age 60-80, density of stocking is reduced to achieve a diameter increment. From 80 years of age, stands reach the felling maturity, forests undergo regeneration, eventually after-regeneration. The rotation period (regulated by law) is in the range of 70 to 170 years for Norway spruce, while in modal sites it fluctuates around 110 years.

The age of tree layer (range: 21-190 years) was classified into 8 age classes of 20-year intervals; the two oldest stands (aged 180 and 190 years) were assigned to the 8th age class. Other variables included in the analysis were: the absolute height yield class (AHYC), *i.e.*, the average height of forest stand at 100 years (range: 14-36 m; modus: 28), reflecting the potential biomass production at a given site; and the density of stocking of forest stands, which is a categorical variable expressed as the ratio between the real stand basal area and the tabular stand basal area (range: 0.3-1). Stands with a value of density of stocking of 0.5 are not represented because of insufficient number of samples. Changes in the density of stocking in production forests are mainly due to targeted management interventions.

Soil sampling and determination of soil organic carbon

Soil samples were taken at the 42 selected monitoring plots during the growing season. Each plot was determined by the FMI as a representative site. Samples were taken from one soil pit (one soil pit at each plot), which was established in the central part of the plot. Before sampling, soil horizons O and A were determined and described. At each pit, one soil sample was taken from the organic (O) horizon and one from the surface mineral (A) horizon

(42 soil samples in O + 42 soil samples in A).

Soil organic carbon (SOC) of samples was determined in a specialized laboratory at the FMI (Brandýs nad Labem, Czech Republic). Before analysis, each sample was homogenized separately for O and A horizons. SOC determination was conducted by the chromo-sulfuric oxidation method, accompanied by the titration of hydroquinone (wet method – Nelson & Sommers 1982). Analysis of samples was performed in two replications. SOC values were converted from percentage units to tonnes per hectare (t ha⁻¹) independently for the forest floor (O horizon) and the surface mineral horizon (A), following to the methodology by Macků (Marek et al. 2011), by considering the thickness of individual horizons at each pit (cm), the bulk density (g cm⁻³), and the SOC content (%). According to the above authors, values of bulk density are different for each FVZ and ecological series, as well as soil horizon, and are part of the methodology.

Statistical analysis

Statistical analysis was conducted using the software STATISTICA[®] ver. 12.0 (StatSoft Inc., Tulsa, OK, USA) to obtain information about the relationships and dependencies among the studied variables. The data included the SOC content in the O and A horizons, the depth of these horizons, humus forms and soil groups, as well as a set of abiotic variables at each plot associated with altitude (mean annual air temperature, mean annual precipitation, duration of the growing season as the sum of days with temperature above 10 °C). Moreover, stand characteristics were included in the analysis (age, density of stocking, absolute height-yield class), as well as site features (edaphic category, forest site, vegetation zonality). The Spearman's correlation coefficient was used to determine the relationships between quantitative variables.

Normality of distribution was verified for each variable using the Shapiro-Wilk test and by visualization of Quantile-Quantile graph (Q-Q graph) and Normal Probability graph (N-P graph). Non-parametric tests were applied where appropriate: the Mann-Whitney U test for the comparison of SOC content between the Podzol and non-Podzol soil groups; the Kruskal-Wallis test followed by a *post-hoc* non-parametric multiple comparison for testing differences in SOC content among plots within different soil groups; the Wilcoxon signed rank test for comparison of SOC content in O and A horizons. Levene's test was used to verify the homogeneity of variance. The homogeneity of the data was additionally evaluated by calculating the coefficient of variation. Outlier identification was carried out using the non-parametric Dean-Dixon test. All tests were performed at 5% level of significance.

Since the SOC content has a non-normal distribution, the median values (50th per-

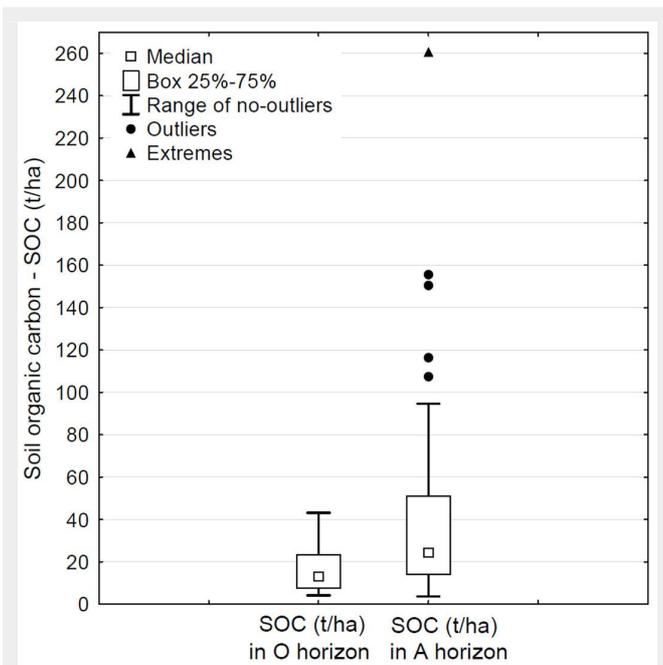


Fig. 1 - Comparison of SOC content ($t\ ha^{-1}$) in the organic (O) and surface mineral (A) soil horizons at the 42 monitoring plots examined in this study.

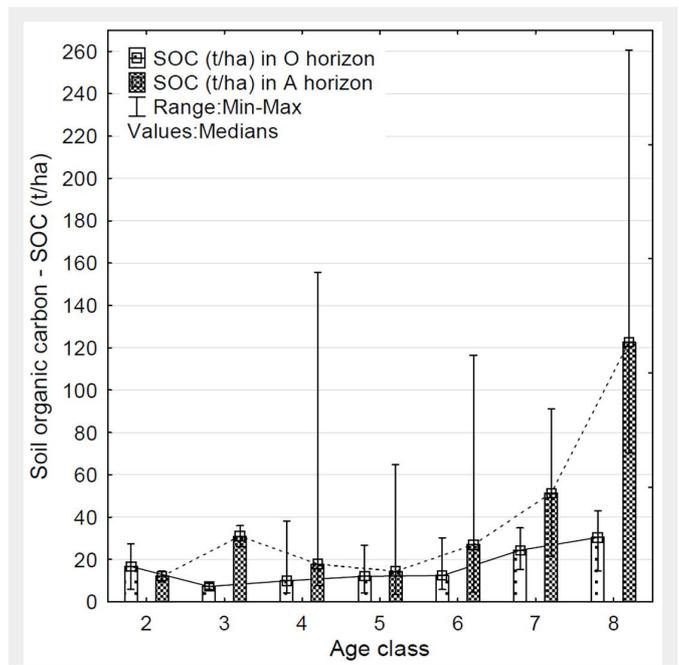


Fig. 2 - Variation of SOC content ($t\ ha^{-1}$) in the organic (O) and surface mineral (A) soil horizons in relation to stand age class in the studied plots.

centile) were used to avoid possible bias of means due to outliers and extreme values. For the same reason, the min-max range is shown in figures; however, standard deviation (SD) and standard error of the mean (SE) for individual age classes and density of stocking are listed in Tab. S3 and Tab. S4 (Supplementary material), to allow a better comparison with median values. Outliers and extreme values were intentionally not removed from the dataset, as they provided one of the findings which are part of the results.

To identify the factors better accounting for SOC variability, linear regression models of log-transformed SOC content were fitted separately for O and A horizons. Firstly, categorical variables were included in the models, and then quantitative variables and their interactions with variables already included in the model were added. Only significant predictors (variables/interactions) were kept in the model. If the interaction between a quantitative and categorical variable was significant, the effect of a quantitative variable was also investigated in individual subgroups defined by the categorical variable.

For better understanding the relationships among quantitative variables, Principal Component Analysis (PCA) was performed and the first two principal components were visualized using a biplot.

Results

The content of SOC in the surface mineral (A) horizon varied within the range of 3.68 to 260.64 $t\ ha^{-1}$ (with a median of 24.38 $t\ ha^{-1}$), with a marked variation due to the presence of several extreme values (outliers – Fig. 1). In the organic (O) horizon the

content of SOC varied within the range of 4.21 to 43.19 $t\ ha^{-1}$ (with a median of 13.11 $t\ ha^{-1}$ – Fig. 1). The content of SOC was higher in the A horizon than in the O horizon in 33 monitoring plots with dominant representation of Norway spruce. In the remaining 9 plots, a higher SOC content was found in the organic horizon. The difference in SOC content between horizons was statistically significant ($p < 0.01$).

A positive correlation between SOC content and stand age class was found for both O and A horizons. In particular, the content of SOC in the organic horizon increases significantly ($R = 0.43$; $p < 0.01$)

with the age class of stands (Fig. 2), with a deviation in the 2nd age class. The same dependence was demonstrated in the A horizon ($R = 0.45$; $p < 0.01$) with deviations in content of SOC in the 3rd and 4th age class (Tab. S3 in Supplementary material). The highest content of SOC was detected in the 8th (oldest) age class of stands in both horizons. On the other hand, the lowest content of SOC in the O horizon was detected in the 3rd age class, while in the A horizon it was in the 2nd and subsequently in the 5th age class of stand. The fourth and fifth age classes in the A horizon are characterized by a decrease in the otherwise in-

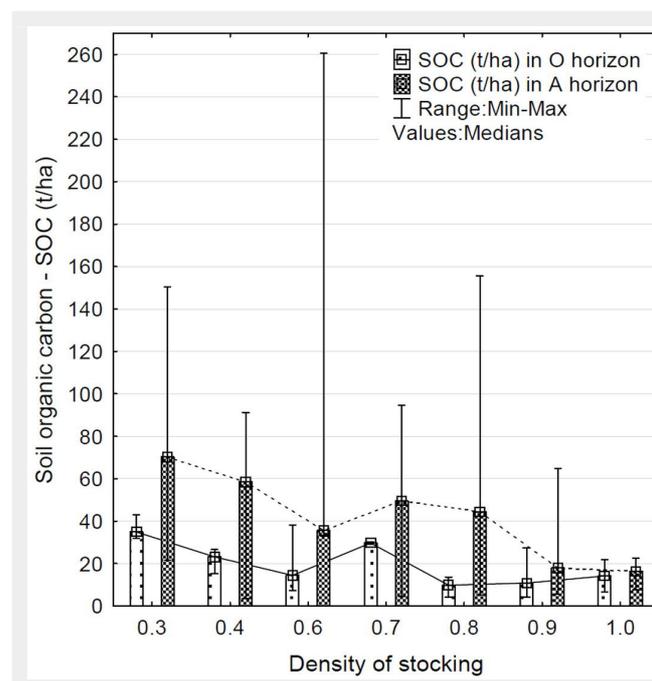


Fig. 3 - Variation of SOC content ($t\ ha^{-1}$) in the organic (O) and surface mineral (A) soil horizons in relation to the density of stocking in the examined forest stands.

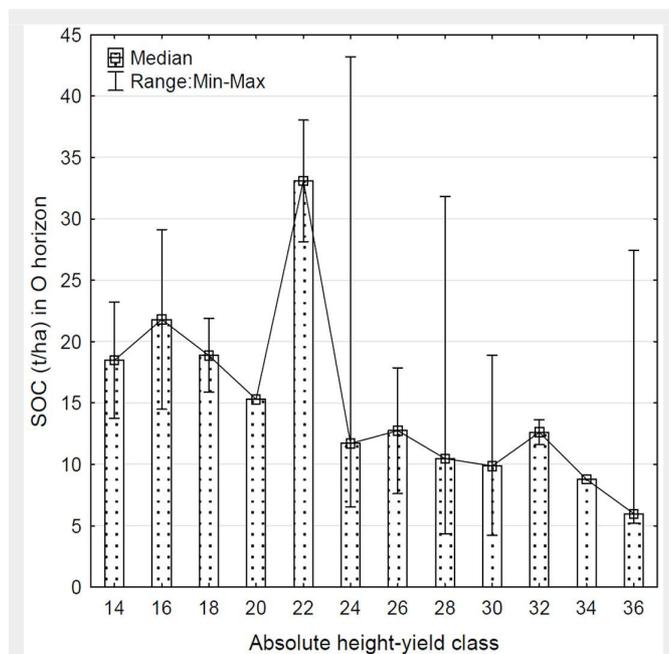


Fig. 4 - Variation of SOC content ($t\ ha^{-1}$) in the organic (O) soil horizons in relation to the absolute height-yield class (AHYC) of stands.

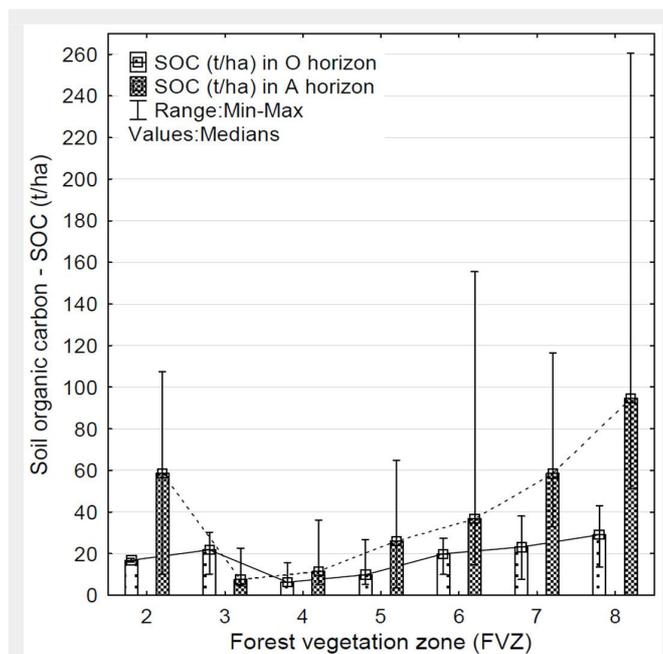


Fig. 5 - Variation of SOC content ($t\ ha^{-1}$) in the organic (O) and surface mineral (A) soil horizons along an elevation gradient. (FVZ): forest vegetation zones (Viewegh et al. 2003).

creasing trend of SOC. Values of SOC content in the A horizon from 4th to 8th age classes had the largest variation. The range of SOC values in the studied horizons according to the age class of stands are shown in Tab. S3 (Supplementary material).

Correlation analysis revealed a negative correlation between soil organic carbon (SOC) and the density of stocking for both horizons (Fig. 3). The content of SOC in the organic horizon decreases ($R = -0.39$; $p < 0.05$) with increasing the density of stocking in the stand. The highest amount of SOC in the O horizon was at a density of stocking of 0.3, i.e., at the lowest values in the selected samples. The lowest amount of SOC in the O horizon was determined at values of 0.8 and 0.9. The correlation value of SOC with the density of stocking in the A horizon was similar ($R = -0.36$; $p < 0.05$) to that of the O horizon. A considerable decline in the trend of SOC content in the A horizon is apparent when density of stocking is 0.6. The lowest level of SOC in the mineral horizon was detected at a density of stocking of 1.0 and 0.9. On the other

hand, the highest level was detected at a density of stocking of 0.3 and 0.4. The content of SOC in the A horizon again showed a remarkable variation (Tab. S4 in Supplementary material) due to extreme values (outliers). The trend of variation in SOC values (median) at increasing density of stocking is shown in Fig. 3.

Analysis of the relationship between age class and density of stocking of stands revealed a negative correlation ($R = -0.52$; $p < 0.01$), i.e., the density of stocking decreases when the age class increases. Up to the 5th age class the value of density of stocking was relatively stable, while a decline occurs in the 6th and 7th age classes; also, 7th and 8th age classes include stands with considerable variation in density of stocking. Distribution of individual values of density of stocking in the age classes of stands in the study area are given in Tab. S2 (Supplementary material).

The relationship between the absolute height-yield class (AHYC) and SOC was also examined in the studied plots. A significant relationship was found, with a negative

correlation between AHYC and SOC only in the O horizon ($R = -0.48$; $p < 0.01$), indicating that when AHYC increases, the SOC content decreases in the organic horizon (Fig. 4), with the exception of AHYC values 16, 22 and 32. No significant correlation was found in the A horizon, though a decreasing trend of SOC in relation to AHYC in the surface mineral horizon was also evident (Fig. S2 in Supplementary material).

A significant relationship was also detected between SOC content and altitude (in terms of FVZ) in the organic ($R = 0.44$; $p < 0.01$) and surface mineral horizons ($R = 0.68$; $p < 0.01$). The content of SOC in the soil increased with increasing FVZ (Fig. 5). In Fig. 5 a gradually increasing trend of SOC is evident from 4th to 8th FVZ for both horizons. A more pronounced increasing trend and a higher SOC content were detected in the A horizon. In the 4th FVZ, 6.34 $t\ ha^{-1}$ of SOC was detected in the O horizon, and 11.45 $t\ ha^{-1}$ of SOC in the A horizon; in the 8th FVZ the value of SOC in the O horizon was 29.14 $t\ ha^{-1}$ and 94.63 $t\ ha^{-1}$ in the A horizon. On the other hand, in the 3rd and 9th FVZ, a higher SOC content was detected in the O horizon as compared with the A horizon. The 2nd FVZ showed a deviation in an otherwise increasing trend of SOC in the A horizon, probably due to the low number of samples and/or the presence of several extreme values (Fig. 5).

SOC content in organic and surface mineral horizons and the altitude were also closely related to the type of soil (soil groups) and humus form (Tab. 1). The highest content of SOC was detected in the Podzol group of soils, which also included a large number of extreme values (Fig. 6). The highest content of SOC was found in

Tab. 1 - Median values of SOC content ($t\ ha^{-1}$) in organic (O) and surface mineral (A) soil horizons, site altitude and prevailing humus form, according to soil group (determined after WRB 2014) represented in the analysed data set.

Soil groups	SOC ($t\ ha^{-1}$)		Altitude (m a.s.l.)		Dominant humus form	Sample size (n)
	O horizon	A horizon	Range	Average		
Haplic Podzols (HP)	28.15	58.52	820-1318	1137	mor	10
Entic Podzols (EP)	13.64	44.46	720-1206	907		8
Cambisols (CA)	10.86	17.21	300-680	534		18
Retisols (RE)	8.40	16.94	390-475	433	moder	2
Stagnosols (SG)	6.66	12.66	420-650	513		4

Haplic Podzols in both horizons (28.15 and 58.52 t ha⁻¹ in O and A horizons, respectively). These samples came from altitudes ranging from 820 to 1318 m a.s.l., where the predominant humus form was mor. Entic Podzols contained 13.64 and 44.46 t ha⁻¹ of SOC in O and A horizons, respectively. These samples came from a range of altitudes from 720 to 1206 m a.s.l., where the dominant humus form was moder. The lowest content of SOC was detected in Stagnosols, with values of 6.66 and 12.66 t ha⁻¹ for O and A horizons, respectively, and moder as prevalent humus form. Although the difference in SOC content among the five soil groups was significant ($p < 0.01$ for both horizons), there was no significant difference between Haplic Podzol and Entic Podzol, and similarly there was no difference among Cambisols, Stagnosols and Retisols. Therefore, we distinguished two soil groups showing significantly different ($p < 0.01$) SOC content: the Podzol and the other (non-Podzol) groups.

SOC content also varied according to the present humus forms (mor and moder). A significantly higher content of SOC was found in samples with mor humus forms in both horizons ($p < 0.01$ – Fig. S3 in Supplementary material). The most pronounced difference in SOC content between mor and moder humus was in the surface mineral (A) horizon (about 40 t ha⁻¹). In the organic (O) horizon a median value of 25.69 t ha⁻¹ was observed, while in the A horizon a median of 64.38 t ha⁻¹ was found for humus form mor.

Log-linear models for SOC content

In both soil horizons, SOC content was significantly higher in the Podzol soil group than in the non-Podzol soil group. The difference in SOC content between mor and moder humus forms was not significant either in Podzols or in non-Podzol soil groups. Previously observed differences between mor and moder humus forms can therefore be explained by a higher frequency of mor humus form in the Podzol soil group: 10 out of 18 Podzols were of mor humus form, whereas only 2 out of the 24 monitoring plots with non-Podzol soil group were of mor humus form.

As expected, all the environmental variables analyzed (mean annual precipitation, mean annual air temperature and duration of the growing season) were highly inter-correlated, as they represent abiotic conditions associated with the elevation gradient. Therefore, only altitude was taken into consideration for inclusion in the log-linear model.

In the Podzol soil group, none of the stand characteristics were significantly associated with SOC content in the organic (O) horizon. In the non-Podzol soil group, none of the variables representing abiotic conditions associated with altitude were significantly associated with SOC content. Neither were associations with age and density of stocking of stands significant.

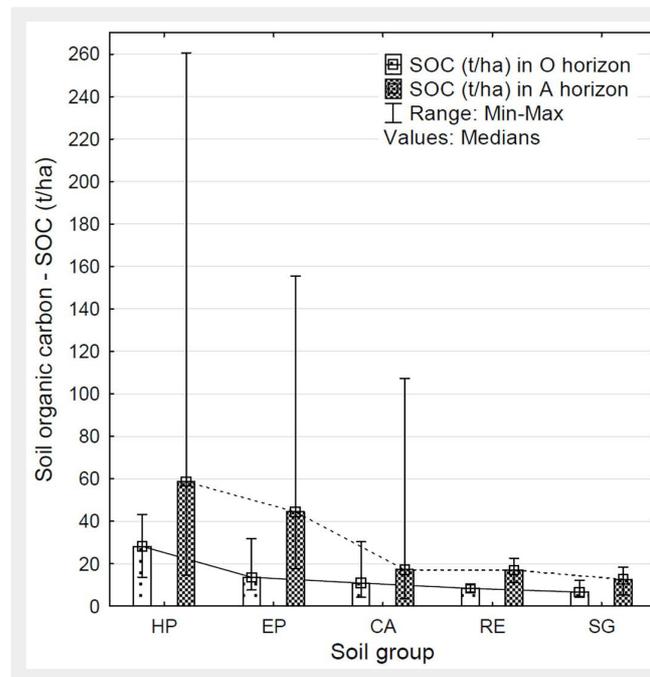


Fig. 6 - Comparison of detected SOC content (t ha⁻¹) in organic (O) and surface mineral (A) soil horizons, according to the soil group (determined after WRB 2014) represented in the analysed data set. (HP): Haplic Podzols; (EP): Entic Podzols; (CA): Cambisols; (RE): Retisols; (SG): Stagnosols.

The only significant predictor of SOC content was AHYC ($p = 0.03$) in the non-Podzol soil group. The association was negative – the higher SOC content in the O horizon was found in stands with low absolute height-yield class.

As regards the SOC content in the surface mineral (A) horizon, no significant association between SOC content and abiotic variables (altitude, mean annual air temperature, mean annual precipitation, duration of growing season) or stand characteristics (age class, density of stocking, AHYC) was found for neither soil groups. In the surface mineral horizon, it was evident that soil group had the most significant influence on SOC content. The effects of other variables were negligible.

The results of the bivariate analyses

showed that soil groups and altitude have a significant influence on SOC content in both horizons. However, when the samples were divided into the 2 groups of soils (Podzols and non-Podzols), the effect of altitude was no longer significant in the surface mineral (A) horizon, while it was significant only for the Podzol soil group in the organic (O) horizon.

Principal components

The PCA revealed the strong correlation among the abiotic variables associated with altitude, which was strongly positively correlated with mean annual precipitation and strongly but negatively correlated with mean annual air temperature and duration of growing season (Fig. 7). These variables had the highest loadings on the first princi-

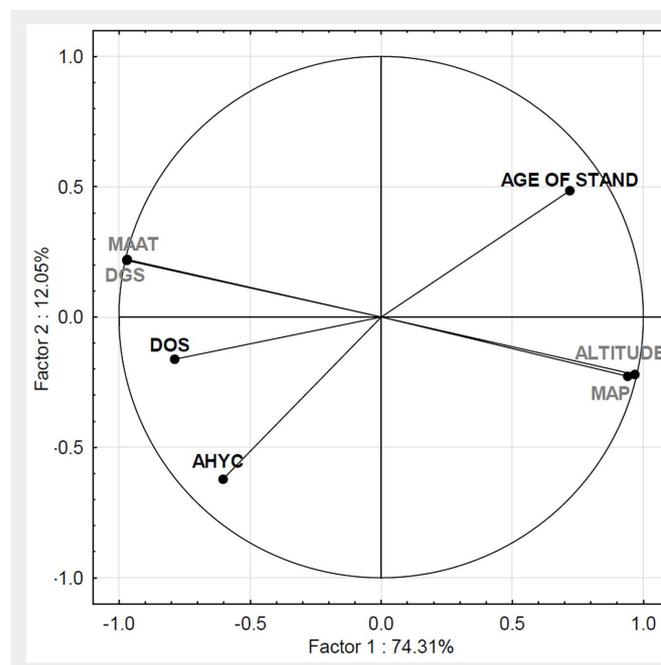


Fig. 7 - Biplot projection of independent variables after PCA. Stand characteristics are marked in black, while variables related to site conditions are in grey. (MAP): mean annual precipitation; (AHYC): absolute height-yield class; (DGS): duration of growing season; (MAAT): mean annual air temperature; (DOS): density of stocking.

pal component, which explained the largest portion (74.3%) of variability, and reflects the abiotic variables associated with the altitude gradient. The second principal component, which explains an additional 12.1% of variability, reflects stand characteristics, and revealed a positive correlation with density of stocking and absolute height-yield class and a negative correlation with stand age.

Discussion

Various studies have highlighted the role of forest ecosystems in the global carbon cycle and in mitigating the impacts of global climate change (IPCC 2000, Janssens et al. 2003, Ciais et al. 2008, Lorenz & Lal 2010, Vesterdal et al. 2013). The potential for carbon storage in the soil has been extensively investigated, mostly focusing on differences among soils beneath different forest tree species (mainly beech and spruce), or by comparing the stock of soil carbon in coniferous and deciduous stands (Oostra et al. 2006, Vesterdal et al. 2008, 2013).

In this work, we studied the potential for carbon sequestration in the soil of 42 representative stands with dominant Norway spruce in the Czech Republic by analyzing their SOC content, with the aim of estimating its amount in the organic and surface mineral soil horizons and assessing the driving factors affecting SOC content. A number of authors (Vesterdal et al. 2013, Pretzsch et al. 2014, Vilén et al. 2015, Andivia et al. 2016, Tashi et al. 2016) agree that the potential for carbon sequestration in the soil is influenced by site conditions (altitude, soil type, soil depth, humus form, etc.) and stands characteristics (age, density of stocking and species composition).

In this study, the relationships between SOC content, abiotic variables, and stand characteristics have been tested by bivariate analysis, log-linear models and PCA. As expected, mean annual air temperature, mean annual precipitation, duration of growing season were highly dependent upon the altitude of the studied stands, represented here by the forest vegetation zones (FVZ). No particular attention has been given to well-known relationships that can be easily explained – for example, the decreasing density of stocking with increasing altitude.

Our results showed that the total SOC stock in forest stands with a dominant proportion of Norway spruce is in the range of 11.91 to 275.11 t ha⁻¹ (with a median value of 37.49 t ha⁻¹ SOC) in all the studied soil profiles (O+A horizon). A significantly higher SOC content was detected in 33 samples of Norway spruce stands in the surface mineral (A) horizon compared to the organic (O) horizon (Fig. 1). This corresponds with the studies by Vesterdal et al. (2008, 2013), who reported that tree species which have a low SOC content in the organic soil horizon have a higher content of SOC in the surface mineral soil horizon. A higher SOC content in the surface mineral horizon was

also observed by previous studies carried out in Norway spruce stands (Gurmesa et al. 2013, James et al. 2014, Andivia et al. 2016). Contrastingly, a higher SOC content in the organic (O) horizon was reported by several authors (Frouz et al. 2009, Harper & Tibbett 2013, Wiesmeier et al. 2013). We observed a higher SOC content in the organic horizon only in 9 monitoring plots out of 42, which can be attributed to the greater thickness of the organic horizon.

Many studies have compared SOC content in the soil organic horizon between different tree species, concluding that conifers in general (*Pinus*, *Tsuga*, *Larix*, *Picea*) and especially Norway spruce (*P. abies*) contain higher SOC stock in the organic horizon than other tree species (Oostra et al. 2006, Vesterdal et al. 2008, Frouz et al. 2009, Laganière et al. 2012, Gurmesa et al. 2013, Vesterdal et al. 2013, Kern et al. 2016). Indeed, the SOC content in this horizon is primarily influenced by species composition (Vesterdal et al. 2008, Gurmesa et al. 2013). Other authors also focused on SOC content in the surface mineral horizon, which is more affected by site conditions (Vesterdal et al. 2008, 2013, Gurmesa et al. 2013).

In the long term, it is preferable that carbon (SOC) is accumulated in the mineral horizon rather than in the organic horizon (Jandl et al. 2007, Frouz et al. 2009, Rumpel & Kögel-Knabner 2011, Vesterdal et al. 2013). In fact, the organic horizon is more vulnerable as it is more influenced by both disturbances (e.g., cutting, fires) and decomposition processes, by which carbon could be released back into the atmosphere. Therefore, in terms of stable carbon sequestration into the soil, the mineral horizon plays a very important role.

In this study, we found that the SOC stock is significantly affected by stand characteristics, in terms of age, density of stocking and absolute height-yield class. An increasing trend in SOC with age class of the stand was detected both for the organic and surface mineral horizons. A similar trend was observed by Hiltbrunner et al. (2013) and Světlík et al. (2016). A slight decrease in the surface mineral horizon was evident in the 4th (61-80 years) and the 5th (81-100 years) age class. The lowest values of SOC were found in stands in the 5th age class. Stands in the 4th and 5th age classes were mostly located at lower altitude, from 2th to 5th FVZ, in sites where forests with a dominance of Norway spruce are not natural. The prevailing humus form on these sites was moder. It is possible to presuppose that this decrease in the trend of SOC content may be caused by relatively better humification compared to higher FVZ. A faster decomposition retroactively releases carbon into the atmosphere as CO₂, which in turn becomes available for potential sequestration of carbon into the tree biomass. The high ability of spruce to allocate carbon into its biomass in these age classes (especially 4th and 5th) has been

previously reported by several authors (Cienciala et al. 2006, Hiltbrunner et al. 2013, Vilén et al. 2015, Parolari & Porporato 2016). This hypothesis has been also confirmed by a long-term research by Vilén et al. (2015), which was conducted in conifer (including Norway spruce) and broad-leaved stands in the Czech Republic and Finland (from 1950-2010). The authors demonstrated that most carbon was allocated in tree biomass, especially in stands in the 4th and 5th age classes. According to Světlík et al. (2016), most carbon in Norway spruce is stored in the stem biomass (about 69.92% of the total carbon of individual tree biomass), while minor proportion are stored in the roots (16.80 %). Grünwald & Bernhofer (2007) reported a carbon sequestration into biomass of 4-7 t ha⁻¹ year⁻¹ in Norway spruce stands at aged 108 years in Germany. However, in coniferous stands with production management, the input of organic matter from tree biomass to forest soil is minimal. Indeed, tree stems with bound carbon are usually exported from such stands for further processing.

Stands of Norway spruce from 6th to 8th age class show a significantly increasing trend of SOC content and a high stock in the organic and surface mineral soil horizons (Fig. 2). In the case of 6th age class, these are trees of felling maturity in production forests, while forest stands in the 7th and 8th age classes have a protective function in the study area, being located at higher altitudes (7th and 8th FVZ), where the dominant humus form is mor. In these sites, the decomposition rate decreases and there is an accumulation of organic matter and carbon in the soil (Drewnik 2006), which could explain the higher SOC content. Based on our results, we can conclude that these forest stands have a manifest potential to sequester carbon into the soil, predominantly in the surface mineral horizon (Carey et al. 2001, Jandl et al. 2007, Luysaert et al. 2008).

The content of SOC and minor differences in both soil horizons in stands in the 2nd age class can be explained by the influence of residual SOC content from the previous (cuttings) generation of forest. An increase in the difference of SOC content in organic and surface mineral horizons in the 3rd age class can be explained by the displacement of the SOC stock through the root systems growing into the deeper layers of soil. Stands in these age classes were represented only by a low number of samples, and were found mainly at lower altitudes.

During the development of forest stands used for production, density of stocking is changed through management interventions, i.e., the density of stocking decreases with increasing age, as confirmed by the significant relationship between stocking density and stand age observed in this study. Stands up to the 5th age class reached almost full stocking (the median value for density of stocking was 0.9),

while a decrease occurs in the 6th age class. The variation of the density of stocking from 4th to 6th age classes relates to the process of forest regeneration. Older protective forest stands in the 7th and 8th age class showed a lower density of stocking, although with considerable variation.

A negative relationship was also found between density of stocking and SOC in both soil horizons. The lowest SOC content in the surface mineral horizon was detected when density of stocking was 0.9–1.0 and stands were mostly in the 4th and 5th age class (Fig. 3). It is possible to conclude that the denser the forest is, the more stable the conditions are for decomposition and humification (Brady & Weil 2002), which reduce SOC content in the soil. Conversely, stands with open canopy with lower values of density of stocking (e.g., 0.3, 0.4 and 0.7) are able to sequester more carbon into the soil, especially into the mineral horizon. Moreover, a significant decreasing trend of SOC with increasing AHYC (absolute height-yield class) was observed only in the organic horizon, suggesting a significant impact of climate. Stands with higher yield class of trees, and thus with lower SOC content in soil, are located mainly in the 3rd to 6th FVZ, where the rate of decomposition and humification is faster, which is evident in the prevailing humus form moder. Stands with lower yield class and higher SOC content are located in the 7th and 8th FVZ (i.e., at sites with a natural occurrence of stands dominated by Norway spruce), where the prevailing humus form is mor, the decomposition processes take place only marginally and the carbon is stored in the soil. As for the AHYC values 16, 22, and 32, which do not conform with the general trend described above (Fig. 4), they include stands occurring on Haplic Podzol and Entic Podzol soil types. As mentioned above, this group of soils is distinguished by a distinctly higher SOC content in both horizons (Tab. 1, Fig. 6), probably due to the humus form present (Tab. 1).

Our results demonstrated a significant influence of altitude and soil group on SOC content in both soil horizons. In particular, the nature of soil (Podzol or non-Podzol) seems to be an important factor that could explain the significant differences in the SOC content observed in both soil horizons. However, stands of the Podzol soil group are located at the highest altitudes (mountain Podzol) in the study area, and are characterized by extreme values of SOC content, therefore outliers were not removed from the dataset. Nonetheless, the increasing trend of SOC content in both soil horizons along the altitudinal gradient (expressed in terms of FVZ – Fig. 5) is consistent with previous findings reported by Garten & Hanson (2006) and Tashi et al. (2016). Moreover, the lowest SOC content in the surface mineral horizon was found for stands located in the 3rd and 4th FVZ (365–580 m a.s.l.), as well as in the 4th and 6th age classes with density of stocking

around 0.9, with AHYC in the range of 18 to 28, where the decomposition process is faster (Fig. 5). These positions are unnatural for Norway spruce in the study area. Based on our results, no significant carbon storage into the soil occurs at these elevations. Furthermore, in the context of regional scenarios of anticipated global climate change (Čermák et al. 2004) and related shift in FVZ, Norway spruce growing outside of its ecological and climatic optimum is expected to suffer of impaired health conditions, reduced growth and ultimately to high mortality (already evident at present). In this context, stands with dominant Norway spruce at the lowest elevation sites are unlikely to reach the 6th age class (around 120 years), when, according to our findings, they begin to show the ability for carbon sequestration into the soil (Carey et al. 2001, Luysaert et al. 2008).

A high and increasing SOC content in both soil horizons was found for stands located from 5th to 8th FVZ (770–1225 m a.s.l.), with the highest SOC content in the 8th FVZ (Fig. 5). There is a slowing of the decomposition process and the accumulation of soil organic matter (Drewnik 2006) due to climatic conditions. It can be stated that in these positions, the most significant carbon sequestration into the soil occurs, especially into the mineral horizon.

Norway spruce forest stands reached the highest yield class in the current 4th, 5th and 6th FVZ in the study area. At higher altitudes AHYC declines, mainly due to the effects of climate. With the expected climate change and the associated shift of FVZ, in terms of process of AHYC, the potential for carbon sequestration into the soil by Norway spruce forests will probably be higher at higher FVZ (i.e., elevation – Fig. S4 in Supplementary material), where Norway spruce naturally occurs and at sites with prevailing Haplic Podzol and Entic Podzol soil types, which showed the highest stock of SOC.

Based on the obtained results, it is apparent that forest stands with a natural dominance of Norway spruce growing at higher altitudes (from 6th to 8th FVZ) in the 6th to 8th age class with lower density of stocking (sparse stands), have a significant ability to sequester carbon into the soil, especially in the surface mineral horizon. Conversely, production forest stands, predominantly in the 4th to 5th age class, at almost full density of stocking (from 0.9 to 1.0), and located at lower altitudes (3rd and 4th FVZ, marginally in the 5th FVZ), showed a lower ability of carbon sequestration into the soil. According to many authors, Norway spruce stands of these age classes tend to demonstrate the ability to fix carbon in their biomass (Cienciala et al. 2006, Vilén et al. 2015, Parolari & Porporato 2016).

Regional scenarios of expected climate change in the study area foresaw a shift of 1–2 forest vegetation zones (Čermák et al. 2004). This change will lead the optimum

conditions for cultivation of Norway spruce to shift up to higher altitudes, i.e., towards the area of its natural distribution (Čermák & Holuša 2010, Hlásny et al. 2011, Hanewinkel et al. 2013). Current Norway spruce stands at lower altitudes and on sites outside its natural habitat will be unsuitable for its further growth and survival (Čermák et al. 2004, Čermák & Holuša 2010, Hlásny et al. 2011, Fitzgerald & Lindner 2013, Hanewinkel et al. 2013). Several studies showed that forest stands with natural dominance of Norway spruce at higher altitudes (> 600 m a.s.l.) can be very important for carbon storage (Janssens et al. 2003, Vetter et al. 2005). Our results suggest that Norway spruce forest stands in their ecological optimum will be capable of long-term survival, sustainable growth and effective carbon sequestration.

Conclusion

Our results show that the studied stand characteristics (age, density of stocking and absolute height-yield class) and site conditions (soil type, humus form, elevation gradient) have an impact on the potential for carbon sequestration in the soil of forest stands dominated by Norway spruce in the eastern part of Czech Republic. A higher content of SOC was demonstrated in the surface mineral soil horizon compared to the organic horizon, which is desirable for the long-term stability of carbon sequestration in the soil.

According to our results, older stands (age: 101–190 years) with lower density of stocking (sparse stands), located in the elevation range 1010–1225 m a.s.l. (6th to 8th FVZ, i.e., the range of natural occurrence of the Norway spruce forest), have the highest potential for long-term stable carbon sequestration in the soil (especially in the mineral soil horizon), even in the context of expected climate change.

Priority adaptive measures should be considered to face the effects of anticipated climate change on carbon sequestration in the soil, including a change of forest management towards near-natural forms which respect the potential natural vegetation. Further research is needed to better assess the effectiveness of individual species (both individually and mixed) in preserving the “carbon service” outside locations and sites of natural representation of Norway spruce.

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Supplementary Material

Tab. S1 - Frequency of the studied plots according to ecological series and edaphic categories.

Tab. S2 - Frequency of the analysed samples from the monitoring plots, according to density of stocking and stand age class.

Tab. S3 - SOC content ($t\ ha^{-1}$) in the organic (O) and surface mineral (A) soil horizon, according to stand age class.

Tab. S4 - SOC content ($t\ ha^{-1}$) in the organic

(O) and surface mineral (A) soil horizon, according to density of stocking of forest stands.

Fig. S1 - Location of the study area and monitoring plots. (1): Dražanská vrchovina Upland; (2): Jeseníky Mountains.

Fig. S2 - Variation of SOC content ($t\ ha^{-1}$) in the surface mineral soil horizon in relation to the absolute height-yield class (AHYC).

Fig. S3 - Comparison of SOC content ($t\ ha^{-1}$) in the organic (O) and surface mineral (A) soil horizon according to the humus forms present in the analysed dataset.

Fig. S4 - Variation of median values of absolute height-yield class (AHYC) in each represented forest vegetation zones (FVZ).

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