

Quantifying forest net primary production: combining eddy flux, inventory and metabolic theory

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Net primary production (N_{PP}) is a central and fundamental carbon-related term in global change studies. We proposed a top-down method to quantifying forest N_{PP} which overcomes the deficits of the traditional bottom-up method. The new top-down method combines eddy flux data, climate variables, tree inventory and metabolic theory. Our method was tested in six forests and provides reliable annual N_{PP} estimations which are consistent with bottom-up results. Carbon use efficiency also supports this new method. Taking advantage of fine temporal resolution of our top-down method, we examined whether and confirmed N_{PP} was well correlated with leaf area index at a seasonal scale, as suggested by past studies. The potential value of our new method as a standard N_{PP} method is high because of the world-wide network on eddy tower and inventory plot, however further data of performance of the new method is needed to fully evaluate its performance under different conditions.

Keywords: Carbon Use Efficiency, Metabolic Theory of Ecology, Autotrophic Respiration, Biomass, Eddy Covariance

Introduction

The earth's surface environment changed dramatically between 1960s and 2010s, including changes in atmospheric gas concentrations, temperatures, and land-cover. Notably, the surface temperature increased by 0.9 °C and the atmospheric CO₂ levels increased by 23% (from 316 to 389 part per million – Brohan et al. 2006). The surface warming was most likely caused by radiative forcing resulting from atmospheric CO₂ buildup. Consequently, understanding the carbon balance of ecosystems is a critical component in global change studies.

Net primary production (N_{PP}) is a central carbon-related term which integrates climatic, ecological, geochemical and human

influences on the biosphere (Nemani et al. 2003). N_{PP} is the rate of biomass change in plants, including woody, leaf and root tissues, but also root exudates and volatile organic carbon compounds.

N_{PP} is defined as the difference between gross primary production (G_{PP}) and total plant respiration (R_a) in an ecosystem (Schulze 2006). In the past, directly measuring both G_{PP} and R_a were impractical due to methodological constraints, especially for tall forest G_{PP} (Clark et al. 2001a). Whole ecosystem R_a was not only difficult to estimate but also involved significant uncertainties (Lavigne 1997). Due to methodological advances, it is now possible to obtain ecosystem G_{PP} by using eddy-covariance techniques (Beer et al. 2010). Furthermore,

a potential way for reliable R_a estimation has been developed based upon metabolic theory (Mori et al. 2010).

The traditional method to quantify N_{PP} was developed during International Biosphere Program (IBP – Newbould 1967), which calculate N_{PP} by combining the organic tissue biomass growth of all quantifiable components within a plant or an area. Currently N_{PP} is normally calculated using the IBP-method with field-surveys (Clark et al. 2001b).

Though widely-used, several intrinsic deficits exist in the IBP-method. First, not all N_{PP} components can be directly measured in the field due to environmental variations during the measuring interval (Clark et al. 2001a). Second, it is difficult to obtain a fine temporal resolution N_{PP} (i.e., daily), and thus to explore the explicit links of N_{PP} to environmental conditions. Third, the IBP-Method is labor intensive and therefore difficult to implement automatically. Methodological constraints not only limit field-data availability, but also hinder the accuracy of N_{PP} estimation and prediction. Therefore, alternative new methods are urgently needed for forest N_{PP} studies.

We propose a new method to quantify N_{PP} , which more accurately quantifies its various components and explores how these components co-vary with climatic variation at a high temporal resolution. The new method was tested in six primary forests. To overcome major deficits of the traditional IBP-method, we directly address weaknesses within the IBP method and compare the results derived from both methods. Top-down design, high temporal resolution, and automatic monitoring represent a significant advance in forest N_{PP} .

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estimation, and could provide a way of comparing N_{PP} of forests globally. It could also serve as an independent way to verify estimation derived from the current IBP-method.

Materials and methods

The new top-down method for N_{PP}

The new method estimates N_{PP} via top-down analysis (eqn. 1):

$$N_{PP} = G_{PP} - R_a$$

The general work flow is illustrated in Fig. 1. Most data required by this method are derived from eddy flux, air temperature, and tree size. Tree size is converted into biomass by using allometric equations. R_a of each tree at 20 °C (R_{20}) was estimated according to the common tree biomass-respiration relationship reported by Mori et al. (2010 – eqn. 2):

$$\frac{1}{R_{20}} = \frac{1}{201.78 M^{1.408}} + \frac{1}{0.410 M^{0.805}}$$

where M is the biomass of a tree. The coefficients were estimated from 271 naturally growing trees in tropical, temperate and boreal forests, with a regression correlation coefficient of 0.99 (Mori et al. 2010). R_a was constructed from the Q_{10} model of Janssens & Pilegaard (2003 – eqn. 3):

$$R_a = R_{ref} - Q_{10}^{\frac{(T - T_{ref})}{10}}$$

where R_{ref} is R_a at the reference temperature T_{ref} , here specified to R_{20} at 20 °C. G_{PP} is obtained by partitioning eddy flux after necessary data treatment.

Forests for testing

Six forests were used to test the new method.

Forest 1: A temperate forest in Changbaisan (CBS), northeastern China. The geographic location is 128° 28' E, 42° 24' N. The CBS forest is primary and over 250 years old with a canopy height of 26 meters. The annual mean temperature was 4 °C and the annual rainfall was 700 mm. The dominant

tree species are *Pinus koraiensis*, *Tilia amurensis*, *Acer mono*, *Fraxinus mandshurica*, and *Quercus mongolica*. Inventory data were collected from a 20 ha plot very close to the eddy flux site (Hao et al. 2008).

Forest 2: A tropical seasonal forest in Xishuangbanna (BNS), Southwest China. The geographic location is 101° 15' E, 21° 55' N. The annual mean temperature and rainfall were 21.7 °C and 1490 mm respectively. The mean canopy height was 30 m. Trees in the topmost layer primarily include *Pometia tomentosa*, *Terminalia myriocarpa*, *Gironiera subaequalis*, and *Garuga floribunda*. Inventory data were collected from a 1 ha plot beneath the eddy flux tower (Zhang et al. 2010).

Forest 3: A tropical seasonal deciduous forest in Mae Klong (MKL), Thailand. The geographic location is 108° 53' E, 18° 43' N. The annual mean temperature and rainfall were 25 °C and 1500 mm respectively. The mean canopy height was 30 m. The stand age was around 30 years in 2008. Domestic species in overstory are *Shorea siamensis*, *Vitex peduncularis*, *Xylia xylocarpa*. Inventory data were collected from a 4 ha plot (Marod et al. 1999).

Forest 4: A tropical lowland rainforest in Pasoh (PSO), Malaysia. The geographic location of the forest is 102° 18' E, 02° 58' N. The annual mean temperature and rainfall were 25.3 °C and 1865 mm respectively. The mean canopy height was 35 m. The studied forest is primary. Domestic species in overstory predominantly belong to Dipterocarpaceae, Leguminosae, Burseraceae. A total of 814 species existed in the 50 ha plot. Inventory data were collected from a 6 ha plot beneath the eddy flux tower (Hoshizaki et al. 2004, Kosugi et al. 2008).

Forest 5: A tropical secondary forest in Bukit Soeharto (BKS), Indonesia. The geographic location is 117° 02' E, 0° 51' S. The annual mean temperature and rainfall were 27 °C and 3300 mm respectively. The studied forest is a secondary forest. In 1998, the secondary-growth forest disappeared due to a forest fire facilitated by El Niño Southern Oscillation (ENSO) event. The mean canopy height was 12.4 m in year 2004. We used eddy flux data in 2001 and 2002 and tree census data in 2003. Domi-

nant species in overstory were *Macaranga gigantea*. Inventory data were collected from three 1 ha plots which were not logged (Ruslim et al. 2000).

Forest 6: A logged tropical rainforest in Tapajos (K83), Brazil. The geographic location is 54° 31' E, 03° 01' S. The annual mean temperature and rainfall were 25 °C and 1911 mm respectively. The mean canopy height was 35 m. The inventory plot is 16 ha in the east of an eddy flux tower (Miller et al. 2004).

The eddy flux based G_{PP}

Eddy covariance based net ecosystem CO_2 exchange (N_{EE}) is estimated as (Monson & Baldocchi 2014 – eqn. 4):

$$N_{EE} = \rho \overline{w'c'} + \rho \int_0^z \frac{dc}{dt} dz$$

The overbar indicates averaging over time, primes represent fluctuations from the average, ρ is the molar density of dry air, w is vertical wind velocity, c is the mixing ratio of CO_2 measured by infrared gas analyzer, dc/dt is the change in CO_2 mixing ratio with time, and z_r is the measurement height.

The N_{EE} could be decomposed into G_{PP} and ecosystem respiration (R_e) as (eqn. 5):

$$G_{PP} = -(N_{EE} - R_e)$$

The minus sign outside the parenthesis results from discipline convention. Total R_e is the sum of nighttime and daytime R_e , and conceptually nighttime N_{EE} is equivalent to nighttime R_e . However, nighttime N_{EE} is usually underestimated during periods of low turbulence. The common way to correct this underestimation is the so-called u^* filtering method. In this method, the threshold u^* was determined by finding a saturation point in u^* and nighttime N_{EE} plot. All N_{EE} observations were filtered by removing all values below the u^* threshold prior to analysis.

Inventory, tree biomass, R_{20} and R_a estimation

Mori et al. (2010) suggests a common tree biomass-respiration relationship. R_{20}

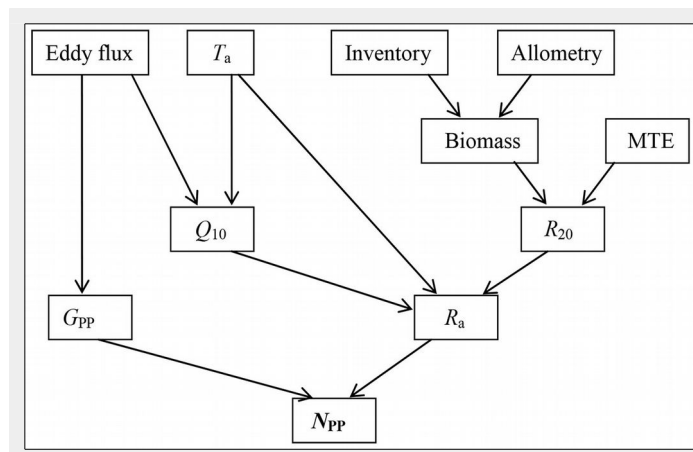


Fig. 1 - Flow schema of top-down estimation of forest net primary production (N_{PP}). Eddy flux is carbon flux measured by eddy covariance technique; (T_a): temperature; (Q_{10}): respiration temperature sensitivity index; (G_{PP}): gross primary production; (R_{20}): tree autotrophic respiration at 20°C; (R_a): autotrophic respiration; (MTE): metabolic theory of ecology here specified to Mori et al. (2010)'s work. Tree census (Inventory) data and site-specific allometric equations were used in calculating tree biomass individually.

was estimated from tree biomass obtained by Kato et al. (1978) for BKS, MKL and PSO; by Chen & Guo (1984) for CBS; by Chambers et al. (2001) for K83; and by Lv et al. (2007) for BNS.

A mixed-power scaling rather than the single-power scaling of eqn. 2 was recommended by Mori et al. (2010). We compared R_{20} estimated with both scaling. On the other hand, we used the Q_{10} derived from flux tower R_e , which includes the temperature sensitivity of both heterotrophic (R_h) and autotrophic respiration (R_a), to obtain R_a with eqn. 3. In some cases, the Q_{10} of R_h and R_a is different and this might lead to biased estimate for R_a . Furthermore, we cannot obtain Q_{10} value for an equatorial tropical rainforest where temperature range was very narrow. Under such circumstance, we used a fixed Q_{10} of 2 for the equatorial rainforest (Amthor 1984, Slot et al. 2014). A sensitivity analysis was also used to investigate how Q_{10} affect R_a .

The IBP based N_{PP}

The detail on estimating N_{PP} with IBP method has been extensively discussed in an IBP book (Newbould 1967) and a review by Clark et al. (2001a). In general, IBP based N_{PP} was estimated as the sum of biomass increment and biomass losses occurred during the investigation interval both above- and below-ground. We did not estimate IBP based N_{PP} for each site but rather collected the published values.

Other supporting measurements and calculations

Remote sensing-derived leaf area index was generally used to estimate N_{PP} at large scales (Gower et al. 1999, Ollinger et al. 2007). We calculated leaf area index (L_{AI}) to examine whether it could explain the seasonal variation of evergreen tropical forest N_{PP} at daily resolution.

Following Jarvis & Leverenz (1983), L_{AI} was estimated as (eqn. 6):

$$L_{AI} = -\ln\left(\frac{Q_T}{Q_0}\right)/\kappa$$

where κ is extinction coefficient which varied between 0.4 and 0.65. We used 0.55. Q_t is the transmitted photosynthetically active radiation (P_{AR}), and was monitored by a sensor (LI-190SB[®], LI-Cor, USA) placed at the height of 3.9m, and Q_0 is the P_{AR} above the canopy measured by a sensor at the height of 36.2 m.

Results and discussion

The R_a derived from metabolic theory and its uncertainties

The application of Mori et al. (2010)'s finding in estimating forest R_a is the key issue of this new method. Mori et al. (2010) claimed that a mixed-power function rather than a single-power function is better in describing the biomass-respiration relationship. The mixed-power function implying a

Tab. 1 - Estimated biomass and reference respiration rate at 20 °C (R_{20}). (MAT): mean annual temperature; (MAP): mean annual precipitation.

| Forest | MAT [°C] | MAP [mm] | Biomass [t dry matter ha ⁻¹] | R_{20} [$\mu\text{mol m}^{-2} \text{s}^{-1}$] | |
|--------------|----------|----------|--|---|-------------|
| | | | | Single power | Mixed power |
| Forest1: CBS | 4 | 700 | 291 | 3.437 | 3.067 |
| Forest2: BNS | 21.7 | 1490 | 375 | 3.998 | 3.545 |
| Forest3: MKL | 25 | 1500 | 265 | 2.416 | 2.090 |
| Forest4: PSO | 25.3 | 1865 | 456 | 4.872 | 4.317 |
| Forest5: BKS | 27 | 3300 | 239 | 2.532 | 2.235 |
| Forest6: K83 | 25 | 1911 | 348 | 4.091 | 3.660 |

gradual ontogenetic transition in the scaling of metabolism is theoretically sound. We compared R_{20} obtained with mixed-power function to that with single-power one (Tab. 1). The single-power based R_{20} is 13% higher than that of mixed-power. In the statistic perspective, both single-power and mixed-power functions have a good ability at simulating R_{20} with correlation coefficient as high as 0.99 (Mori et al. 2010). The R_{20} , however, is very sensitive to the choice of single- or mixed-power function.

Other than R_{20} , another important parameter which affects R_a estimation is the Q_{10} value. Currently, data for Q_{10} of R_a are scarce. As shown in the flow chart of this new method (Fig. 1), we used Q_{10} of R_e derived from eddy flux observations. However, neither R_e nor soil respiration Q_{10} of equatorial tropical forest is available partly due to narrow temperature ranges. Some studies even found a negative or quadratic relationship between soil respiration and temperature (Hanpattanakit et al. 2015). Furthermore, R_e is the sum of autotrophic (R_a) and heterotrophic components (R_h). The Q_{10} derived from flux tower R_e , which includes the temperature sensitivity of both R_a and R_h , might induce some bias when Q_{10} of R_a and R_h is not the same.

Concerning Q_{10} , we analyzed the sensitivity of Q_{10} with R_a (Fig. 2). When the Q_{10} value varied from 1.4 to 2.5, R_a changed slightly for sites which have mean annual tempera-

tures near 20 °C, (i.e., BNS – Fig. 2). However, these changes could be up to around 40% for other sites (Fig. 2). This means the Q_{10} sensitivity is a critical issue in the new method. The Q_{10} value for R_a is available for CBS and BNS sites and it was used to drive the new method. For the other five tropical forests where Q_{10} is not available, it was set as 2.0 (Amthor 1984, Slot et al. 2014) and the corresponding R_a estimates are shown in Tab. 2.

The third relevant issue which may induce additional uncertainties is the use of temperature response function to evaluate R_a . Several alternative functions are available for this case (Gomez-Casanovas et al. 2013). For example, the highly cited Lloyd & Taylor (1994)'s function as (eqn. 7):

$$R_a = R_{ref} \exp\left[E_0 \left(\frac{1}{T_{ref} - T_0} - \frac{1}{T - T_0}\right)\right]$$

where E_0 and T_0 are two parameters which could be related to Q_{10} as (eqn. 8):

$$Q_{10} = \exp\left[\frac{10 E_0}{(T - T_0)(T + 10 - T_0)}\right]$$

Similar to this, the classic Arrhenius function could be related to Q_{10} as (eqn. 9):

$$Q_{10} = \exp\left[\frac{10 E_a}{RT(T + 10)}\right]$$

where E_a is the activation energy. We com-

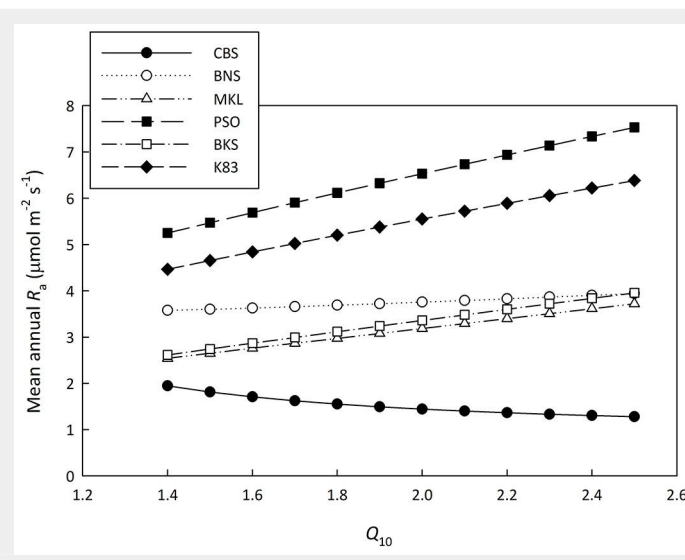


Fig. 2 - Sensitivity analysis of R_a to temperature sensitivity index (Q_{10}). The mean annual autotrophic respiration rate (R_a) estimated with Q_{10} varied from 1.4 to 2.5 for each site.

Tab. 2 - A summary table for eddy flux based gross primary production (G_{PP}), IBP based net primary production ($IBP-N_{PP}$), metabolic theory based autotrophic respiration (R_a) and the new top-down method based N_{PP} (New- N_{PP}). All values are in the unit of $tC\ ha^{-1}\ yr^{-1}$. (‡): For CBS and BNS, Q_{10} was derived from eddy flux based R_{eT} , for other forests Q_{10} was set at 2.0; (§): values in parentheses represent the range from estimates when $Q_{10}=1.4$ to $Q_{10}=2.5$; (£): values after “±” in this column are 20% uncertainties, as suggested by Desai et al. (2008); other values after “±” in this table are standard deviations (STD); (1) Chen et al. (2013); (2) Zhang et al. (2010); (3) Dlioksumpun et al. (2009); (4) Kosugi et al. (2008); (5) Hirata et al. (2008); (6) Malhi et al. (2009); (7) Wang et al. (2006); (8) Zhang et al. (2010); (9) Kira (1978); (10) Toma et al. (2000)

| Forest | Eddy flux G_{PP} | IBP- N_{PP} | R_a ‡ | New- N_{PP} £ |
|--------------|--------------------|------------------|----------------------|-----------------|
| Forest1: CBS | 13.38 (1) | 7.69 (7) | 5.32 ± 0.31 | 8.06 ± 1.61 |
| Forest2: BNS | 23.42 ± 1.74 (2) | 8.80 (8) | 14.63 ± 0.86 | 8.79 ± 1.75 |
| Forest3: MKL | 23.98 (3) | 7.97 (3) | 12.06 (9.60-14.08) § | 11.92 ± 2.38 |
| Forest4: PSO | 32.16 ± 1.32 (4) | 12.85 (9) | 24.72 (19.87-28.50) | 7.44 ± 1.48 |
| Forest5: BKS | 27.10 (5) | 14.70 (10) | 12.71 (9.87-14.97) | 14.39 ± 2.87 |
| Forest6: K83 | 31.40 (6) | 14.40 ± 1.30 (6) | 21.00 (16.90-24.15) | 10.40 ± 2.08 |
| Mean ± STD | 25.24 ± 6.85 | 11.07 ± 3.27 | 15.07 ± 6.90 | 10.17 ± 2.64 |

pared the application of Lloyd & Taylor (1994)’s function to the Q_{10} function of eqn. 3 with the dataset in CBS site. The uncertainty caused by this difference is small. To simplify the methodology, we eventually used eqn. 3 for all cases in this study.

Eddy flux based G_{PP} and its uncertainties

One crucial component of the new top-

down method is reliably calculating G_{PP} with eddy covariance based carbon fluxes (Fig. 1). The G_{PP} can be obtained by decomposing N_{EE} as shown in eqn. 5. Extrapolating the post u^* filtering nighttime R_e to that of daytime is a common method within the flux community (Reichstein et al. 2005). It has successfully been used to reproduce global G_{PP} (Beer et al. 2010).

Nevertheless, the Reichstein et al. (2005) method was criticized for not providing a reliable G_{PP} value. Leaf respiration may be inhibited under light conditions. Using values extrapolated from nighttime, but rather that of daytime, overestimates R_e and subsequently the G_{PP} (Wohlfahrt et al. 2005a, Wehr et al. 2016).

As an alternative method, daytime respiration could be also obtained by extending daytime measurements to zero light, commonly termed as light response method (Wohlfahrt et al. 2005b, Lasslop et al. 2010). Rather than an extensive discussion and comparison on decomposition techniques separating N_{EE} into G_{PP} and R_e , we used CBS site’s data as a case analysis. The light response regression was done with a window of 10 days by using the Michaelis-Menten function as (eqn. 10):

$$N_{EE} = - \left(\frac{\alpha P_m I}{\alpha I + P_m} - R_d \right)$$

where I is the light intensity, α is the apparent quantum yield representing the slope under weak light, P_m is the light-saturated photosynthesis rate, R_d is the derived daytime ecosystem respiration. The temperature dependence of R_d and nighttime R_e of CBS site is shown in Fig. 3. Nighttime R_e was generally higher than that of daytime R_e derived from light response. The R_e estimation was compared between that derived from nighttime data only and that based on both day- and nighttime data (Fig. 4). This supports the claim that R_e with nighttime flux was overestimated (Wehr et al. 2016), though the methods include some difficulties to quantify components.

(i) Can R_d accurately represent the daytime ecosystem respiration? R_d estimate is very sensitive to weak light observations. Weak light intensity usually occurs under transitional periods in the early morning, late afternoon or in periods of high cloud cover. Rapid and large changes in atmospheric stability will cause unnecessary uncertainties in this process, i.e., the flux burst effect in the morning (Grace et al. 1996).

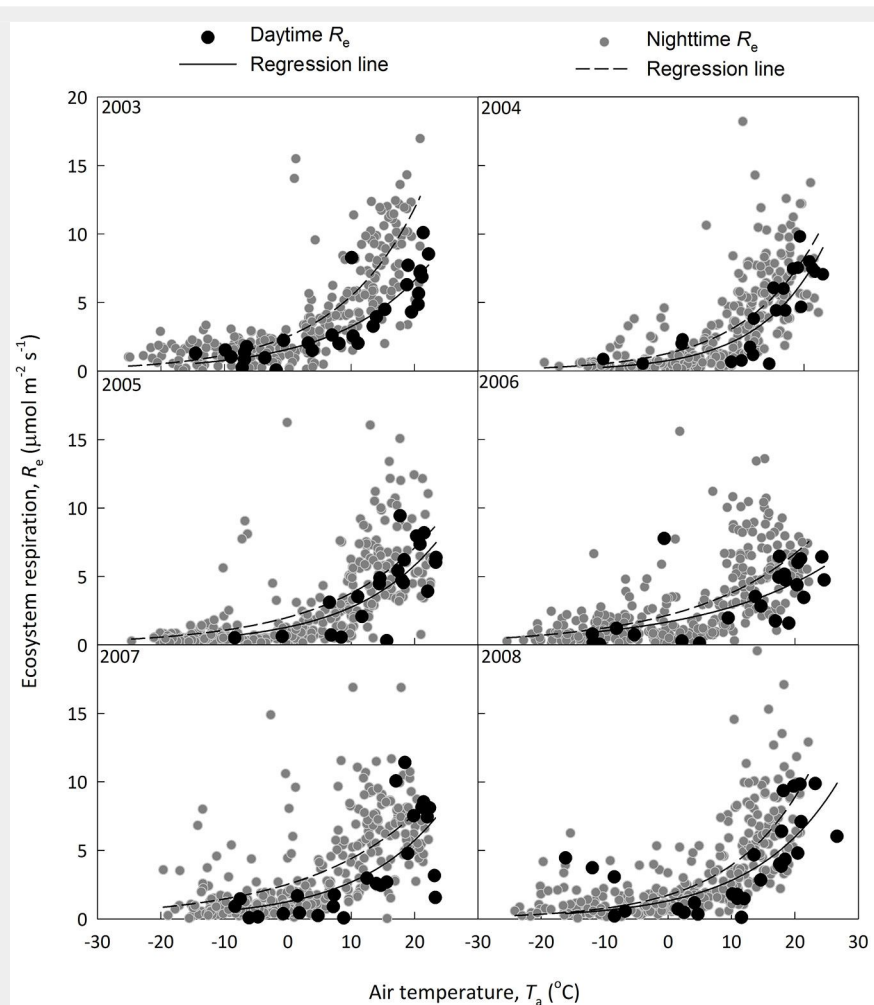


Fig. 3 - A comparison on daytime ecosystem respiration (R_e , black circles and solid regression line) derived from light response and nighttime R_e (dark grey circles with dashed line). The Q_{10} model (see text for detail) was used in fitting the dataset.

(ii) Could R_d derived from the simple light response model represent the daytime ecosystem respiration in reality? A group of variables drive daytime fluxes including R_d . Lasslop et al. (2010), for example, showed that a model water vapor pressure deficit (VPD) included could work better in obtaining R_e and G_{PP} than a model drive by temperature only.

Uncertainties along with decomposition techniques in deriving G_{PP} are substantial. Desai et al. (2008) compared 23 decomposition methods and reported that most methods differed up to 20% in estimates of both G_{PP} and R_e (see also Reichstein et al. 2005, Wohlfahrt et al. 2005b, Desai et al. 2008, Lasslop et al. 2010, Reichstein et al. 2012).

In this study, we cite published G_{PP} values for each site (Tab. 2) as a defensible, persuasive and unified decomposition technique on eddy flux based G_{PP} is currently unavailable (Desai et al. 2008, Wehr et al. 2016). Eddy flux data processing usually needs further site-specific information. We also compare the reported G_{PP} to that obtained with an online processing procedure (<http://www.bgc-jena.mpg.de/bgii/index.php/Services/REddyProcWeb>) widely used in European Flux communities for CBS and BNS, both of which show high levels of agreement. Thus, the up-to-date G_{PP} reported by site's investigators could be viewed as a good option.

The matching of inventory plot with eddy flux footprint

To apply the proposed method, we should ensure that the plot used for the inventory is comparable with the eddy covariance flux footprint for that site. Nevertheless, the scale of inventory plot and eddy flux footprint is usually not well matched both in space and time. While an inventory plot is usually set in a few hectares, an eddy flux footprint could be as large as $\sim 1 \text{ km}^2$. In the current study, we selected six forests to validate the proposed method. The inventory plots of CBS, BNS, and PSO sites were just located within the eddy flux footprint. For the K83 site, inventory and eddy flux data have already been used in a matching method to address carbon balance (Miller et al. 2004). The forest type and disturbance history were the same in the inventory plot and eddy footprint for all six sites.

Another helpful experimental design is setting multiple small plots instead of a big one. Multiple small plots distributed in the footprint, when all footprint biomass estimation is not affordable, are definitely better in terms of representativeness than a single big one placed somewhere in the vicinity (Law et al. 2008). This helpful design was not involved in the current study but was strongly recommended for future studies.

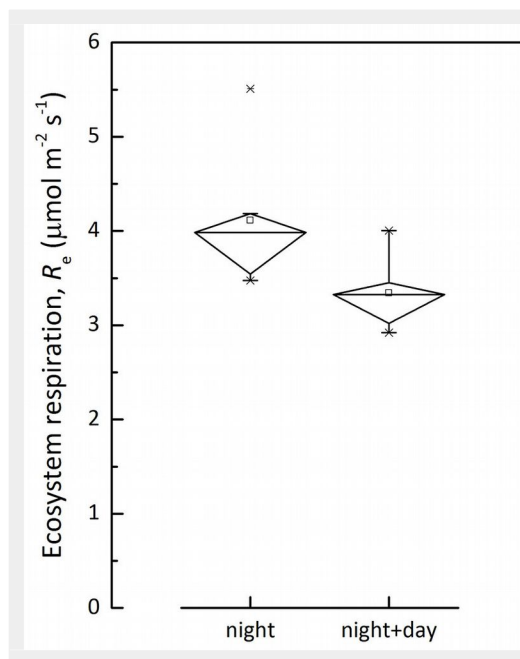


Fig. 4 - A comparison on mean ecosystem respiration rate estimated with nighttime data only (night) and a combining of both day- and night-time data (night+day).

Comparison on the new top-down N_{PP} and IBP based N_{PP}

Both the new top-down N_{PP} and the IBP based N_{PP} estimates are listed in Tab. 2, and a paired Student *t*-test showed no significant difference ($p=0.55$ – Fig. 5). We also found N_{PP} estimates are nearly identical for the CBS and BNS sites where Q_{10} is available, providing strong support for the new top-down method.

Nevertheless, the new N_{PP} differed from the IBP N_{PP} in the MKL, PSO and K83 sites. For the PSO and K83 sites, the new N_{PP} seems to be underestimated whilst R_a overestimated. This underestimation may be related to the assumed fixed Q_{10} value. Conversely the MKL site shows the opposite, and the biomass of the MKL site is as low as that of the temperate CBS forest.

The MKL forest is under its secondary growth. The stand age of this secondary forest is only around 38 years and the inventory data used to calculate biomass and R_a was collected in year 1992, when the stand age was only 14. Since biomass is expected to increase during the early stages of secondary growth, the biomass and R_a derived with data collected in 1992 is likely to be underestimated, which may account for the overestimation of R_a with the new method in MKL.

Further support for the new top-down methods can be derived from the carbon use efficiency (C_{UE}), *i.e.*, the ratio between N_{PP} and G_{PP} . C_{UE} was reported to be about 0.50 in temperate forests (Waring et al. 1998, DeLucia et al. 2007, Zanotelli et al. 2013) and about 0.35 in tropical forests

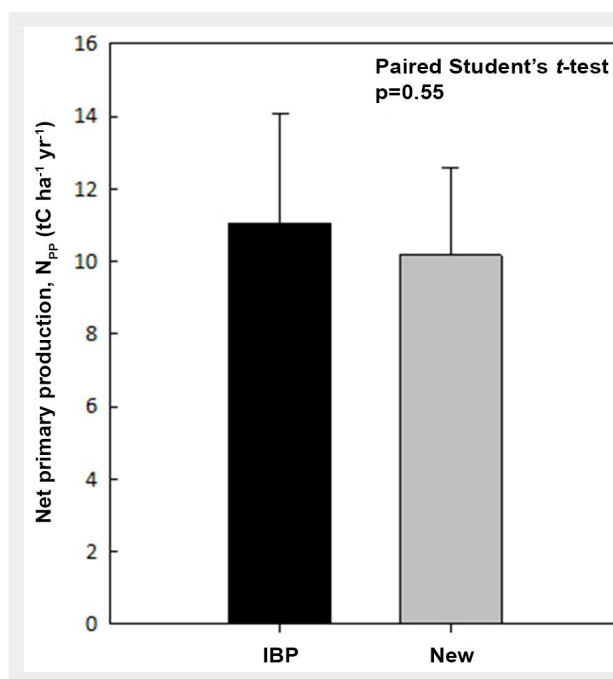


Fig. 5 - A comparison on net primary production (N_{PP}) estimated with IBP method (IBP-NPP) and the new top-down method (New-NPP).

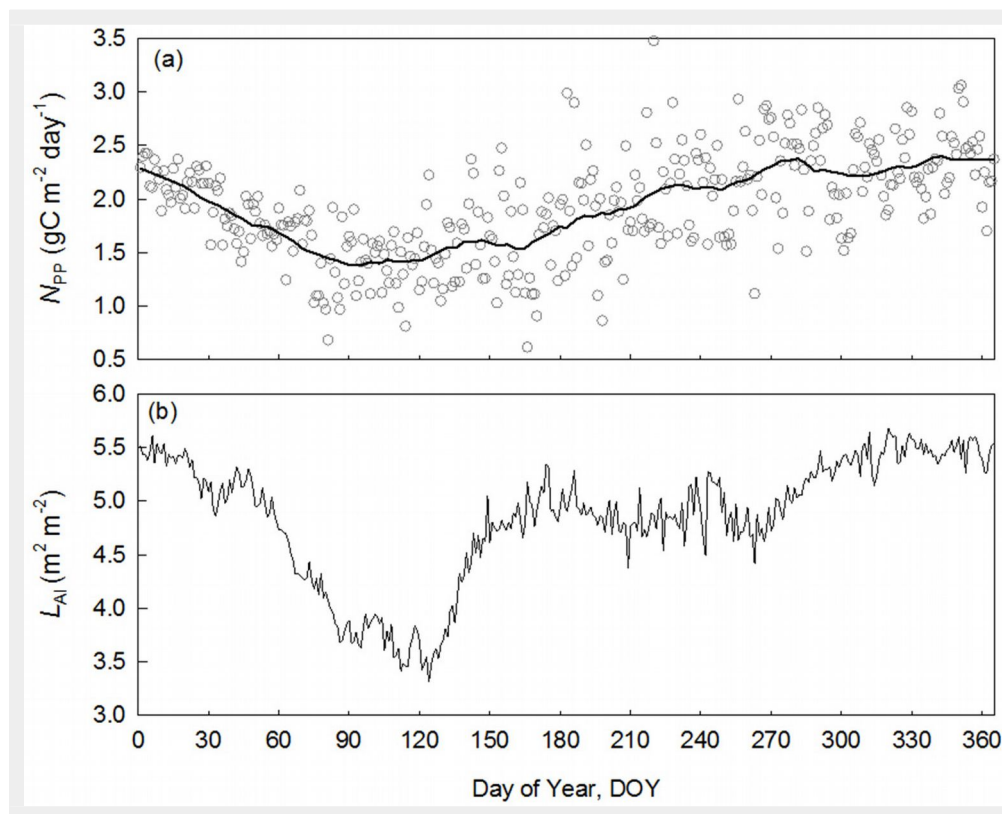


Fig. 6 - Seasonal variation of net primary production (N_{PP}) calculated by the proposed new method and leaf area index in a tropical rainforest of Xishuangbanna (BNS), China. (a): daily N_{PP} in open circles in black grey, solid line show smoothing line; (b): daily leaf area index (L_{AI}) derived from light transmittance.

(Malhi 2012). C_{UE} derived from the new top-down method was 0.60 in temperate forest and around 0.3 for primary tropical forests (C_{UE} is 0.37, 0.23, 0.33 for BNS, PSO, K83, respectively), which agree well with the previous suggestions.

The seasonal variation of N_{PP}

Accurately quantifying the seasonal variation of forest N_{PP} by the IBP method is challenging. The tradition temporal resolution for IBP based N_{PP} is a year (Clark et al. 2001a), which ignores temporal variability in climatically variable sites. Most of our knowledge of N_{PP} seasonality is indirectly obtained through remote sensing (Gower et al. 1999, Ollinger et al. 2007). The remote sensing inversion of N_{PP} is based on the strong relationship between N_{PP} and the vegetation index. We utilized the finest temporal resolution available for the top-down N_{PP} to test the relationship between the N_{PP} and the vegetation index.

The seasonality in N_{PP} is clear in BNS (Fig. 6a), and is lowest during the late dry season when water vapor pressure deficit (V_{PD}) and soil water deficit peak (indicated by soil water content, S_{WC}). The leaf area index (L_{AI}) was also lowest in the period. A close relationship was found between L_{AI} and N_{PP} (Pearson's $r=0.56$, $n=365$). This result confirmed the reliability of remote sensing derived N_{PP} and provides a direct estimate on seasonal N_{PP} .

Currently, many forest sites have installed eddy covariance systems and established inventory plots. A world-wide test on this new top-down method will both advance our knowledge on N_{PP} and give a more solid terrestrial N_{PP} estimation. The limita-

tions and uncertainties of the new top-down method are substantial. Incorporating other process based models into the top-down method would increase accuracy and may enhance its value as a standard method.

Conclusions

Forest N_{PP} is a fundamental parameter describing the ecosystem functioning. It is the change rate of biomass including woody, leaf and woody tissues, exudates and volatile organic carbon compounds. We proposed a novel method to quantify N_{PP} using measurements that are different from the traditional IBP method. We propose the use of eddy covariance together with the total biomass and the scaling factor proposed by Mori et al. (2010) to estimate the autotrophic respiration (R_a). Then N_{PP} is simply obtained from the gross primary production (G_{PP}) derived from eddy covariance, with the simple equation $N_{PP} = G_{PP} - R_a$.

Six sites were used to validate the proposed new method. They were selected because of the good match between inventory plot and eddy flux footprint. The new method provides reliable annual N_{PP} estimations which are consistent with bottom-up results. This consistence was also supported by the past knowledge on carbon use efficiency. Taking the advantage of fine temporal resolution of the new top-down method, the seasonal changes of N_{PP} at daily interval was presented. A close relationship between N_{PP} and leaf area index was found at seasonal scale. This finding validated the new method on the one hand, and provided further support of

using satellite vegetation index to address N_{PP} dynamics on the other. The potential of this new method as a standard N_{PP} method is high because of the world-wide network on eddy tower and inventory plot, however further work of using the method with cases under environmental stress is needed to fully evaluate its performance under variable conditions.

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