A comparison between stomatal ozone uptake and AOT40 of deciduous trees in Japan

Hoshika Y, Shimizu Y, Omasa K

A comparison of the maps of stomatal ozone uptake (AF\textsubscript{0}) and concentrations exceeding 40 ppb (AOT40) for dominant temperate deciduous tree species (Quercus serrata, Fagus crenata, Betula ermanii) was conducted in Japan. Estimations of stomatal ozone uptake were accomplished using estimated ozone concentration, climate data, and vegetation data. Key parameters such as stomatal conductance parameters for each species were collected from scientific literature in Japan. Stomatal closure induced by vapour pressure deficit affected the AF\textsubscript{0} values in warmer part of Japan. For this reason, the areas with high AOT40 did not always correspond to the areas with high AF\textsubscript{0}. The result showed that ozone risk assessment using AOT40 is VPD-constrained in central Japan, which implies an overestimation of risk compared to AF\textsubscript{0}. While in Europe AOT40 is higher where water stress is recurrent, AOT40 peaked in the cool and humid climate region of central-eastern Japan where also stomatal ozone uptake reached maximum values.

Keywords: Deciduous forest trees, Ozone uptake modeling, Stomatal conductance, Japan, Ozone

Introduction
Phytotoxic nature of ozone has been well known for decades (e.g., Nies 1980, Nies 1984). Surface ozone concentrations are increasing in East Asia because of rapid increases in emissions of the main critical ozone precursors, Nitrogen oxides and volatile organic compounds (Naja & Akimoto 2004). Ohara & Sakata (2003) reported that the annual average concentration of photochemical oxidant, mainly ozone, increased with high rate (0.33 ppb year\textsuperscript{-1}) from 1985 to 1999 in Japan. Therefore, assessments of ozone impacts on plants have become very significant in Japan.

In Japan, effects of ozone on the growth of forest tree species have been investigated by using AOT40, which represents the cumulative exposure above a hourly threshold concentration of 40 ppb during daylight hours (e.g., Iizuka et al. 2001, Kohno et al. 2005). Kohno et al. (2005) suggested that 8-15 ppm·h AOT40 for sensitive species (e.g., Fagus crenata), and 15-30 ppm·h AOT40 for moderately sensitive species (e.g., Quercus serrata) from April to September as critical level for forest species, are able to induce a 10% growth reduction. Critical levels for forest trees are set at 5 ppm·h AOT40 corresponding to a 5% growth reduction in Europe (UNEP 2004). An assessment approach using AOT40 has the advantage of being simple because only atmospheric ozone concentration data are needed. However, ozone damage to plants depends not only on atmospheric ozone concentrations but also on stomatal ozone uptake into leaves. Several studies suggested that stomatal ozone uptake was closely related to ozone damages (e.g., Omasa et al. 2002, Paolletti & Manning 2007). For this reason, a stomatal flux-based approach using stomatal ozone uptake is expected to provide better assessments of ozone damage to plants in Europe (Emerson et al. 2000). Critical levels for ozone risk using the stomatal flux-based indices were set for potato and wheat, and provisionally for sensitive forest species (beech, birch - UNECE 2004). Karlsson et al. (2007) reported that the stomatal flux-based indices were superior to AOT40 for ozone-sensitive species, based on a reanalysis of published ozone exposure-response data.

Very little has been done to develop the stomatal flux-based approach for assessing ozone damage to plants in Japan. The assessment approach using AOT40 was suggested to be appropriate for Japan because of much precipitation and limited water stress in Japan (Watanabe et al. 2010). However, this assumption has not been verified yet. In this study, we developed a model of cumulative stomatal ozone uptake (AF\textsubscript{0}) and assess the spatial distribution of AF\textsubscript{0} compared to AOT40 for deciduous forest tree species in Japan.

Materials and methods

Vegetation data
Deciduous forest tree species were classified into three forest types within two climatic zones (cool-temperate and warm-temperate zones) in Japan (Nakashizuka & Iida 1995). Cool-temperate zone and warm-temperate zone are mainly in northern Japan and western Japan, respectively. Regarding the forest types, cool-temperate mixed broadleaf/conifer forest and cool-temperate deciduous forest are distributed in the cool-temperate zone. Cool-temperate deciduous forest is just in central-eastern Japan, with a humid climate throughout the year and heavy snow in winter. In the cool-temperate deciduous forest, beech (Fagus crenata) is the dominant species. The cool-temperate mixed forest is dominated by deciduous oak (e.g., Quercus crispula) and birch (Betula ermanii). In the warm-temperate zone, the forest type is classified as warm-temperate deciduous forest, and is dominated by deciduous oak (e.g., Quercus serrata). Three dominant species (Quercus serrata; Fagus crenata; Betula ermanii) were selected for comparison between AF\textsubscript{0} and AOT40. The distributions of the species targeted by this study are shown in Fig. 1. A distribution was determined based on the vegetation data of the National Survey on the Natural Environment, investigated by the Ministry of the Environment (http://www.biodic.go.jp/J-IBIS.html). Spatial resolution of these data was 1 × 1 km.

Input data
Climate data (air temperature, air humidity, solar radiation and wind speed) and ozone concentration data for the year 2000 (Takigawa et al. 2007) were input into our model. These data were estimated using CHASER (Sudo et al. 2002) and the WRF/Chem model (Grell et al. 2005) provided at 6h temporal resolution, a spatial resolution of 40 × 40 km across Japan. We used these data as input data near ground surface at 20 m high. Validation for the model estimates of ozone concentration was conducted at several monitoring sites in Ja-
pan (Takigawa et al. 2007). Although further validation and model improvement may be needed, the model estimates were used for estimation of stomatal ozone uptake.

Input data of distributions for deciduous tree species were derived from identifying whether the grid space (40 × 40 km) included the distribution of vegetation data shown in Fig. 1 or not. The estimation of stomatal ozone uptake was conducted in each grid square (40 × 40 km) with 6 h time steps.

**Estimation of AOT40**

AOT40 was estimated using ozone concentration at the top of the canopy as recommended for ozone risk assessment (UNECE 2004). We used ozone concentration data at 20 m high as C (z_i). AOT40 was calculated as follows (eqn. 1):

$$AOT40 = \int_{t_1}^{t_n} \max\{C(z_i)-40.0\} \, dt$$

where C (z_i) is averaged ozone concentration for 6 hours at the top of the canopy (ppb), n is the number of data for ozone concentrations. AOT40 was calculated over daylight hours with a solar irradiation higher than 50 W m⁻² from April to September (e.g., Kohno et al. 2005).

**Stomatal conductance model**

Leaf-level stomatal conductance of water vapor (g_sw) was estimated using the multiplicative model (Jarvis 1976, Emberson et al. 2000 - eqn. 2):

$$g_{SW} = g_{max} \cdot f_{phen} \cdot f_{light} \cdot \max\{f_{temp} \cdot f_{VPD} \cdot f_{SMD}\}$$

where g_{max} is the maximum stomatal conductance. The other functions are limiting factors of g_{max} and are scaled from 0 to 1. f_{min} is the minimum stomatal conductance and is set to 0 in this study because we could not get data of g_{min} from the literatures. f_{max} is the variation in stomatal conductance with leaf age, and f_{light}, f_{temp}, f_{VPD}, and f_{SMD} are functions of photosynthetically photon flux density at the leaf surface (PPFD, μmol photons m⁻² s⁻¹), temperature (T, °C), vapor pressure deficit (VPD, kPa), and volumetric soil water content (θ, m⁻³ m⁻³), respectively.

The variation in stomatal conductance with leaf age (f_{max}) modifies g_{max} as a function of time within the leaf duration (Emberson et al. 2000). Maruyama & Honda (1993) reported that it took about 1 month for g_{sw} to reach its peak from leaf onset, and to decline due to its senescence until leaf fall for Japanese beech and deciduous oak species. Therefore, we assumed that f_{max} increased linearly from 0 to 1 during the first 30 days after leaf onset and decreased linearly from 1 to 0 during the 30 days prior to leaf fall.

The functions (f_{light}, f_{temp} and f_{VPD}) have been expressed in various forms (e.g., Jarvis 1976,
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Sirisampan et al. 2003, Emberson et al. 2000). In this study, the following formulas were selected because they have been frequently used in modeling studies in Japan (e.g., Sirisampan et al. 2003). These functions are expressed as follows (eqn. 3, eqn. 4, eqn. 5):

\[
\begin{align*}
  f_{\text{light}}(0.5) &= \frac{\text{PPFD}}{\text{PPFD} + f_{\text{light}}(0.5)} \\
  f_{\text{temp}} &= \frac{\text{T} - \text{T}_{\text{min}}}{\text{T}_{\text{opt}} - \text{T}_{\text{min}}} \left( \frac{\text{T}_{\text{max}} - \text{T}_{\text{opt}}}{\text{T}_{\text{max}} - \text{T}_{\text{min}}} \right) \\
  f_{\text{VPD}}(0.5) &= \left[ 1 + \left( \frac{\text{VPD}}{f_{\text{VPD}}(0.5)} \right)^a \right]^{-1}
\end{align*}
\]

where \( f_{\text{light}}(0.5) \) is the value of PPFD when \( f_{\text{light}} = 0.5 \), and \( \text{T}_{\text{opt}}, \text{T}_{\text{min}}, \) and \( \text{T}_{\text{max}} \) represent the optimum, minimum, and maximum temperatures for stomatal conductance, respectively. \( f_{\text{VPD}}(0.5) \) is the value of VPD when \( f_{\text{VPD}} = 0.5 \), and \( a \) is a constant.

Tab. 1 lists the parameters used in the stomatal conductance model. Parameters for \( f_{\text{light}}, f_{\text{temp}}, f_{\text{VPD}}, \) and \( g_{\text{max}} \) were determined from a review of scientific literature on stomatal conductance of temperate deciduous forest tree species in Japan: *Quercus serrata* (Tanaka et al. 1998, Sirisampan et al. 2003, Yamazaki et al. 2006), *Fagus crenata* (Iio et al. 2004), *Betula ermanii* (Muraoka & Koizumi 2005, Yamazaki et al. 2006). Fig. 2 shows a plot of these functions in the stomatal conductance model.

Parameters of \( f_{\text{SMD}} \) could not be included in the model because of insufficient published data about \( f_{\text{SMD}} \) for temperate deciduous tree species in Japan. The \( f_{\text{SMD}} \) values were not calculated in this study (\( f_{\text{SMD}}=1 \)), although this assumption may lead to an overestimation of stomatal ozone uptake. However, Sirisampan et al. 2003 reported that the soil water content had no effect on the stomatal conductance for six tree species in central Japan.

![Fig. 2 - Plots of the functions \( f_{\text{light}}(A); f_{\text{temp}}(B); f_{\text{VPD}}(C) \) of the stomatal conductance model for three deciduous tree species. \( f_{\text{light}}, f_{\text{temp}}, \) and \( f_{\text{VPD}} \) are functions of photosynthetically photon flux density at the leaf surface (PPFD, \( \mu\text{mol photons m}^{-2} \text{s}^{-1} \)), temperature (T, °C), and vapor pressure deficit (VPD, kPa), respectively.](http://www.sisef.it/iforest/130)

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Japan. Therefore, the results would still allow a comparison of the maps between stomatal ozone uptake and AOT40.

Estimation of stomatal ozone uptake

Estimation of stomatal ozone uptake (\(F_{st}\); nmol O\(_2\) m\(^{-2}\) s\(^{-1}\)) were calculated based on the assumption that the ozone concentration at the top of the canopy ([C (z)]) represented a concentration near sunlit leaves at the top of the canopy (Emberson et al. 2007). Then stomatal ozone uptake was calculated as follows (eqn. 6):

\[
F_{st} = C(z) \frac{1}{(r_c + r_s)} \left( g_{sw}/1.65 \right) \left( g_{c}/1.65 + g_{sw} \right)
\]

where \(r_c\) is the leaf boundary layer resistance (s m\(^{-1}\)), \(r_s\) is the leaf surface resistance [1/ (\(g_{sw}/1.65 + g_{ce}\); s m\(^{-1}\))], 1.65 accounts for the difference in diffusivity of water in air compared with ozone, and \(g_{ce}\) is the external leaf or cuticular conductance (s m\(^{-1}\); \(g_{sw}\) was set to 0.0004 s m\(^{-1}\) (Emberson et al. 2007).

Leaf boundary layer resistance (\(r_s\)) was calculated from the wind speed at canopy height, \(u (z)\) in m s\(^{-1}\), and the mean leaf width, \(L_s\) (m, UNECE 2004 - eqn. 7):

\[
r_s = 1.3 \times 150 \left( \frac{L_s}{u(z)} \right)^{0.5}
\]

where the factor 1.3 accounts for differences in diffusivity between heat and ozone. \(L_s\) for oak and beech/birch were determined as 0.1 and 0.05 m, respectively, from leaf shape data in Japan (Endo 1940).

In Europe, the accumulative stomatal ozone uptake (\(AF_{st}\), Y) was recommended to assess the ozone risk for forest species (Karlsson et al. 2007). It is given by the following equation (eqn. 8):

\[
AF_{st} = \int_{Y}^{\infty} \left( F_{st} - Y \right) dt
\]

where \(Y\) is a threshold of stomatal ozone uptake (nmol O\(_2\) m\(^{-2}\) s\(^{-1}\)). In Europe, a threshold \(Y\) is currently used (Emberson et al. 2007, Karlsson et al. 2007). However, it is not clear whether the same threshold of \(F_{st}\) can be applied in Japan. Therefore, we did not set a threshold for the \(F_{st}\) value (\(AF_{st}\)) in the present study.

Estimation of leaf duration

Leaf duration was simply assumed based on phenological data in Japan (Tab. 2). We chose the data by including several species in Japan. Because we could not collect enough phenological data for each species, we assumed no difference in leaf duration among species. From these data, a leaf onset time was set depending on the latitude in Japan (leaf onset time = 3.4 + latitude - 9.1; \(r^2 = 0.89\)). In contrast, time of leaf fall did not show a linear relationship with latitude.

### Table 2 - Observed time (day of year) of leaf onset and leaf fall for deciduous forest trees in Japan.

<table>
<thead>
<tr>
<th>Region</th>
<th>Observation (day of year)</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Leaf onset</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Western Japan (35°10’ N 132°40’E)</td>
<td>107</td>
<td>Ozaki et al. (2000)</td>
</tr>
<tr>
<td>Central-eastern Japan (34°55’ N 137°45’E)</td>
<td>109</td>
<td>Fujimoto (2008)</td>
</tr>
<tr>
<td>Central-eastern Japan (36°30’ N 138°20’E)</td>
<td>121</td>
<td>Kato &amp; Hayashi (2008)</td>
</tr>
<tr>
<td>Northern Japan (39°90° N 141°00’E)</td>
<td>131</td>
<td>Aoki &amp; Hashimoto (1995)</td>
</tr>
<tr>
<td>Hokkaido (42°59’ N 141°23’E)</td>
<td>135</td>
<td></td>
</tr>
<tr>
<td>Leaf fall</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Western Japan (35°10’ N 132°40’E)</td>
<td>339</td>
<td></td>
</tr>
<tr>
<td>Central-eastern Japan (34°55’ N 137°45’E)</td>
<td>319</td>
<td></td>
</tr>
<tr>
<td>Central-eastern Japan (36°30’ N 138°20’E)</td>
<td>297</td>
<td></td>
</tr>
<tr>
<td>Northern Japan (39°90° N 141°00’E)</td>
<td>298</td>
<td></td>
</tr>
<tr>
<td>Hokkaido (42°59’ N 141°23’E)</td>
<td>304</td>
<td></td>
</tr>
</tbody>
</table>

### Table 3 - Mean values (± SD) of estimated AOT40 and AF,0 for deciduous tree species in each region of Japan. The values were calculated for each region including more than 5 grids. Within the same parameter, values with different letters are significantly different (Kruskal-Wallis test, \(p < 0.05\) - number of 40 × 40 km grids, \(p < 0.05\)).

<table>
<thead>
<tr>
<th>Region</th>
<th>AOT40 (ppm·h)</th>
<th>AF,0 (mmol m(^{-2}))</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Hokkaido (42°-44°N 140°-145°E)</td>
<td>5.1 ± 1.6 a</td>
<td>26.7 ± 1.6 a</td>
<td>31.9 ± 2.3 a</td>
</tr>
<tr>
<td>Eastern Japan (35°-42°N 139°-142°E)</td>
<td>16.7 ± 5.0 b</td>
<td>35.1 ± 3.2 b</td>
<td>40.8 ± 4.4 b</td>
</tr>
<tr>
<td>Central-eastern Japan (34°-38°N 136°-139°E)</td>
<td>28.0 ± 5.6 c</td>
<td>41.1 ± 2.7 c</td>
<td>48.3 ± 4.0 c</td>
</tr>
<tr>
<td>Central Japan (33°-36°N 133°-136°E)</td>
<td>30.8 ± 3.5 d</td>
<td>40.1 ± 2.1 c</td>
<td>-</td>
</tr>
<tr>
<td>Western Japan (31°-35°N 130°-133°E)</td>
<td>26.4 ± 4.3 e</td>
<td>41.8 ± 2.2 c</td>
<td>-</td>
</tr>
</tbody>
</table>

### Significance of difference

Kruskal-Wallis test < 0.001 < 0.001 < 0.001 < 0.001

Fig. 3 - A map of estimated AOT40 values from April to September in Japan along the 40 × 40 km grid of the CHASER and WRF/Chem models used to estimate climate and ozone input data.
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Therefore, the averaged value of the time (DOY = 311, s.d. = 17) was adopted as time of leaf fall in this study.

Statistics and data analysis

To compare AOT40 and AF$_{0}$, maps of both indices were provided. In addition, averaged AOT40 and AF$_{0}$ were calculated for each region. The linear correlations between both indices, and between f$_{VPD}$ in summer and AF$_{0}$ were tested. Moreover, difference of the indices among regions was tested with a Kruskal-Wallis test. Results were considered significant at p < 0.05. Statistical analysis was performed with SPSS software version 11.5 (SPSS, Chicago, USA).

Results

Fig. 3 shows a map of the AOT40 values from April to September in Japan, and Tab. 3 shows an averaged value of AOT40 and AF$_{0}$ of each region in Japan. There were differences in the AOT40 values among regions. The AOT40 values showed the highest value more than 30 ppm·h from central-eastern Japan (34°-38°N 136°-139°E) to western Japan (31°-35°N 130°-133°E - Fig. 3). In contrast, the AOT40 values were relatively low (0-10 ppm·h) in Hokkaido (42°-44°N 140°-145°E).

Fig. 4 (A-C) shows maps of the AF$_{0}$ values of the deciduous tree species in Japan. For Quercus serrata, AF$_{0}$ showed the highest value in central-eastern Japan (Fig. 4A). The averaged AF$_{0}$ values in this region were 41.2 mmol m$^{-2}$ s$^{-1}$ (Tab. 3). In this area, the AF$_{0}$ values reached about 50 mmol m$^{-2}$ for Quercus serrata and Fagus crenata, and 55 mmol m$^{-2}$ for Betula ermanii. In Hokkaido or eastern Japan (35°-42°N 139°-142°E), the AF$_{0}$ values were lower than in central-eastern Japan (Fig. 4, Tab. 3). In central Japan (33°-36°N 133°-136°E), the AF$_{0}$ values were similar to the levels in eastern Japan for Quercus serrata (Fig. 4A), Tab. 3). The AF$_{0}$ values showed below 30 mmol m$^{-2}$ for Quercus serrata in this region (Fig. 4A).

A comparison between the AOT40 and AF$_{0}$ maps (Fig. 3, Fig. 4) suggested a difference in the spatial pattern. Although the AOT40 values were different between eastern Japan, and central Japan, the AF$_{0}$ values were not significantly different for Quercus serrata (Tab. 3). These areas are the warmer parts of Japan. Fig. 5 shows the averaged f$_{VPD}$ values for Quercus serrata in summer (July-August). In central Japan, the averaged f$_{VPD}$ values were about 0.8 in the summer (Fig. 5) because of high temperature. The lower f$_{VPD}$ values corresponded with the lower AF$_{0}$ areas for Quercus serrata.

For further analysis, we conducted a correlation analysis between both indices in Japan (Tab. 4). The result shows that the correlation between AOT40 and AF$_{0}$ was lower for Quercus serrata compared to other two species. Quercus serrata is mainly distributed in warm-temperate climate region (Fig. 1). Especially, the correlation was not significant in warmer part of Japan such as central and western Japan (Tab. 4). This result was influenced by effects of VPD on stomatal conductance (Fig. 5, Tab. 5).

For the other two species distributed in humid and cool region and high mountainous area, the correlation between AOT40 and AF$_{0}$ was very high (Tab. 4). Fagus crenata is also distributed in the warmer part of Japan such as central and western Japan (Fig. 1).
Fig. 5 - Spatial distribution of averaged f_VPD values for Quercus serrata from July to August in Japan. f_VPD shows the response of stomatal conductance to vapor pressure deficit, scaled from 0 to 1.

Tab. 4 - Correlation between AOT40 and AF_0 (accumulative stomatal ozone uptake) for each tree species in Japan. The correlation coefficient was calculated for each area including more than 5 grids. (*) and (**) denotes the significance at 5% and 1 % level, respectively. (NS) indicates no significant correlation.

<table>
<thead>
<tr>
<th>Region</th>
<th>Quercus serrata</th>
<th>Fagus crenata</th>
<th>Betula ermanii</th>
</tr>
</thead>
<tbody>
<tr>
<td>All regions</td>
<td>0.33**</td>
<td>0.89**</td>
<td>0.91**</td>
</tr>
<tr>
<td>Hokkaido - Central-eastern Japan (34°-44°N 136°-145°E)</td>
<td>0.53**</td>
<td>0.92**</td>
<td>0.93**</td>
</tr>
<tr>
<td>Central Japan - Western Japan (31°-36°N 130°-136°E)</td>
<td>-0.06 NS</td>
<td>0.34*</td>
<td>-</td>
</tr>
</tbody>
</table>

Tab. 5 - Correlation between f_SMD (response of stomatal conductance to vapor pressure deficit) in summer and AF_0 (accumulative stomatal ozone uptake) for Quercus serrata in Japan. The correlation coefficient was calculated for each area including more than 5 grids. (*) and (**) denotes the significance at 5% and 1 % level, respectively. (NS) indicates no significant correlation.

<table>
<thead>
<tr>
<th>Region</th>
<th>Value</th>
</tr>
</thead>
<tbody>
<tr>
<td>All regions</td>
<td>0.43**</td>
</tr>
<tr>
<td>Hokkaido - Central-eastern Japan (34°-44°N 136°-145°E)</td>
<td>0.38**</td>
</tr>
<tr>
<td>Central Japan - Western Japan (31°-36°N 130°-136°E)</td>
<td>0.9**</td>
</tr>
</tbody>
</table>

the warmer part of Japan, the correlation between AOT40 and AF_0 decreased for Fagus crenata (r = 0.37).

Discussion

The purpose of this study was to compare the spatial map of the AF_0 and AOT40 values for deciduous forest trees in Japan. Emberson et al. (2000) and Simpson et al. (2007) similarly compared the spatial distribution of the stomatal ozone uptake and AOT40 values in Europe. The AOT40 value was relatively high in southern Europe compared to northern Europe. The spatial gradient of the AOT40 values was remarkable from south to north, but the spatial gradient of the stomatal ozone uptake was much less. Stomatal ozone uptake was limited due to drought stress during the summer in Mediterranean regions where ozone concentrations were high (Paoletti 2006).

In this study, AF_0 showed similarly much less regional differences than AOT40. However, there is a difference of the result compared to Europe. In Europe, the exposure-based critical level for forest species is suggested. Values of the critical levels are defined as 5 ppm·h of AOT40. AOT40 values exceeding the European critical levels were shown in 89% of Japan. In Japan, Kohno et al. (2005) suggested that 15-30 ppm·h AOT40 for moderately sensitive spe-

cies (e.g., Quercus serrata) from April to September as critical level for forest species. AOT40 values exceeding 30 ppm·h were shown in 21% of Japan. Some regions where the AOT40 values reached 30 ppm·h corresponded to cool and humid climate such as central-eastern Japan. Averaged stomatal ozone uptake was estimated to be higher than 40 mmol·m⁻² for the three species in central-eastern Japan (Tab. 3). In contrast, stomatal ozone uptake for Quercus serrata did not show any difference between eastern and central Japan, although AOT40 in central Japan was twice that in eastern Japan. Emberson et al. (2000) showed that VPD played a major role in limiting the stomatal ozone uptake. Also in this study, VPD is a limiting factor of the stomatal ozone uptake especially in warmer part of Japan (Tab. 5). For this reason, this region showed a discrepancy between the AF_0 values of Quercus serrata and AOT40. These results suggest that not only ozone concentration but also stomatal closure induced by VPD affected the AF_0 in the warmer part of Japan. The results of correlation analysis for Fagus crenata and Betula ermanii suggested that AF_0 and AOT40 values are similar for ozone risk assessment in humid and cool climate region, because the correlation between AOT40 and AF_0 was very high (Tab. 4), i.e., humid and cool climate was favorable for stomatal opening. However, in the warmer part of Japan, the correlation between AOT40 and AF_0 decreased for Fagus crenata (r = 0.37), because VPD induced a stomatal closure.

In warmer part of Japan, we also could find the difference of spatial pattern of the AF_0 values between Fagus crenata and Quercus serrata (Fig. 4). Stomatal conductance for Fagus crenata was less sensitive to high VPD (>2 kPa) than that for Quercus serrata (Tab. 1; Fig. 2). Therefore, the result is due to the species-specific characteristics of the stomatal conductance parameters. The assessment approach using AOT40 does not consider the difference of species-specific stomatal response to the climate factors. The result of this study illustrated the limitations of performing ozone risk assessment using AOT40 in Japan.

In this study, f_SMD could not be applied in the model because of insufficient published data about soil moisture deficit. However, Sirisampan et al. (2003) reported that the soil water content had no effect on the stomatal conductance of six tree species in central Japan. However, soil water deficit may be a limiting factor in estimating AF_0 even for short periods during a growing season. The parameter of f_SMD would bring about a further refinement of the calculations for Japan. In addition, we assumed that plants absorbed ozone continuously when climate condition was favorable for stomatal ope-
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However, when leaves experience a longer period of ozone exposure, stomatal response would be changed (Paoletti & Grulke 2005). Pleijel et al. (2002) and Danielsson et al. (2003) reported that ozone exposure over a long time caused a decline in gs in crops for their modelling studies. This effect has not included in the model of this study yet. Long-term ozone exposure would also be limiting factor of stomatal ozone uptake. Including the effects of ozone on stomatal conductance would be expected to improve the ozone uptake modelling for deciduous forests in Japan.

Conclusions

The maps between AF0 and AOT40 for dominant temperate deciduous tree species in Japan were compared. Estimations of stomatal ozone uptake were accomplished using estimated ozone concentration and climate data, and vegetation data. As a result, stomatal closure induced by vapor pressure deficit affected the AF0 values in warmer parts of Japan. Therefore, the areas with the highest AOT40 did not correspond to the areas with the highest AF0, suggesting that the use of AOT40 for protecting Japanese forests may be misleading.

Stomatal ozone uptake was suggested for assessing ozone damage more than three decades ago in Japan (Omasa et al. 1979). Many studies (e.g., Paoletti & Manning 2007) suggested that the stomatal flux-based approach is scientifically-sound and would be a useful tool for ozone risk assessment. This paper is a contribution to develop the stomatal flux-based approach in Japan. Future work should validate and improve our model with field measurements. Parameterizations for more species would also contribute further developments because ozone sensitivity of plants and stomatal response to environmental condition are species-specific (e.g., Omasa et al. 2000).

Acknowledgements

This work was supported by JSPS Fellowships for Young Scientists and was partly supported by the Global Environment Research Fund (C-062) of the Ministry of the Environment, Japan.

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